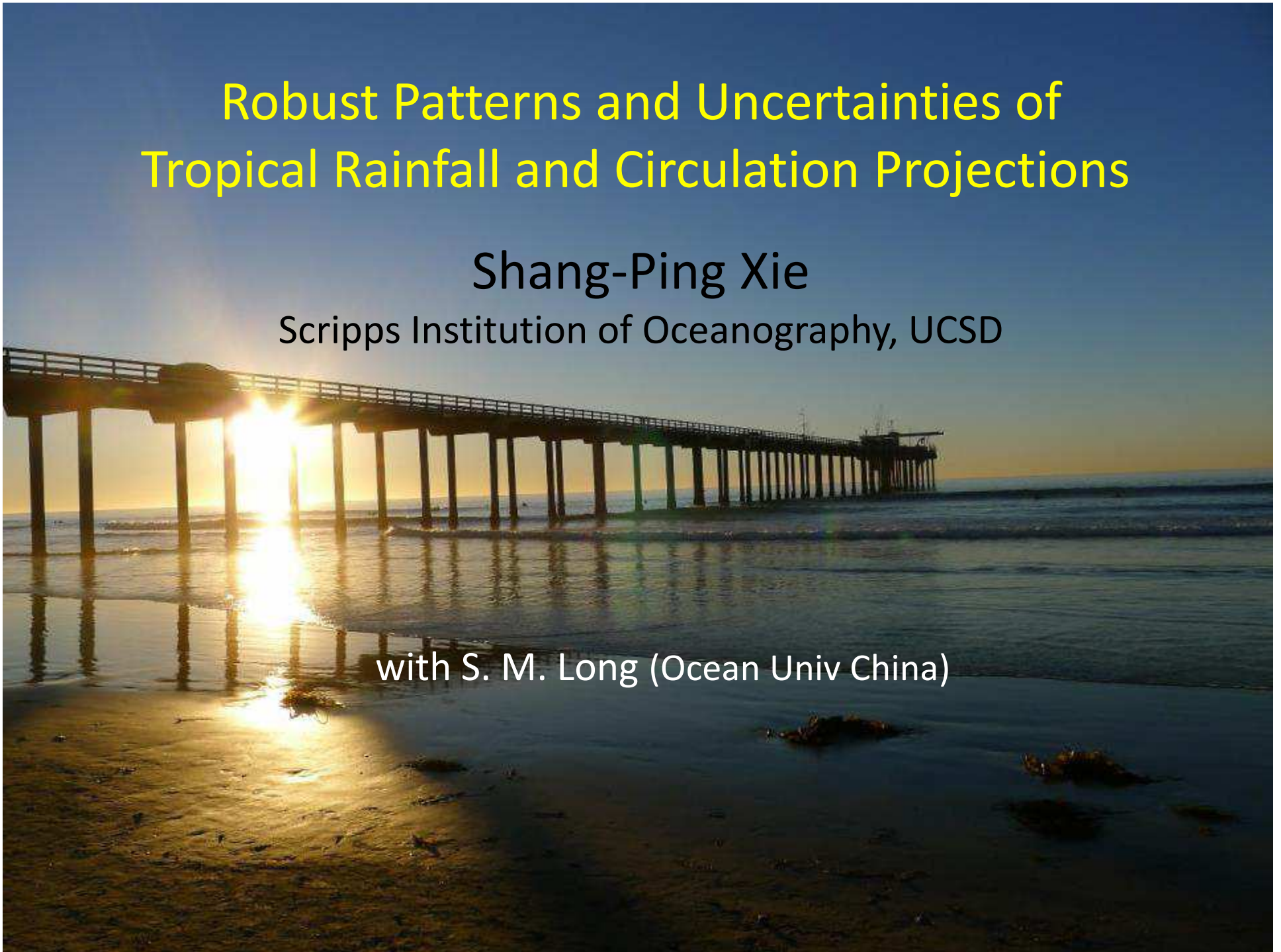


Robust Patterns and Uncertainties of Tropical Rainfall and Circulation Projections

Shang-Ping Xie

Scripps Institution of Oceanography, UCSD

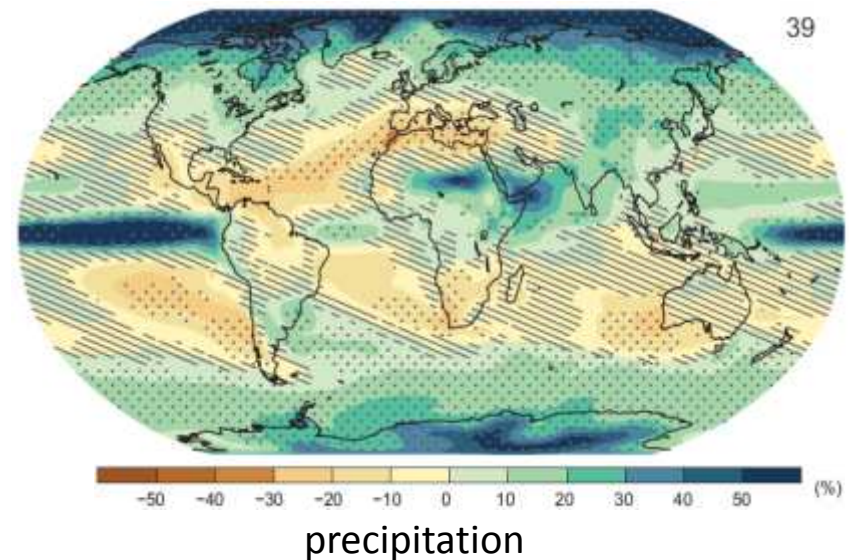
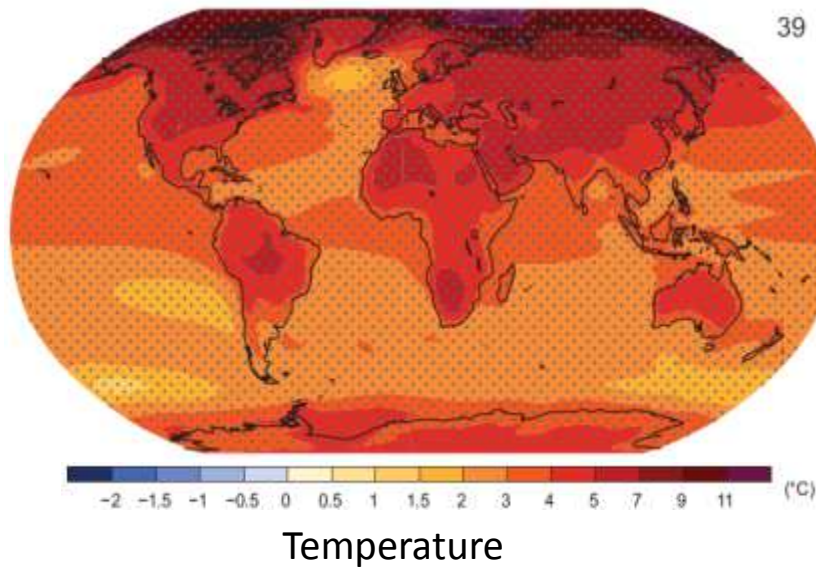
with S. M. Long (Ocean Univ China)



Precipitation change is to first order spatially variable.

- What determines patterns of rainfall change?
- Can we predict the pattern?

Climate change (1986–2005 to 2081–2100), Business as usual (RCP8.5); IPCC AR5

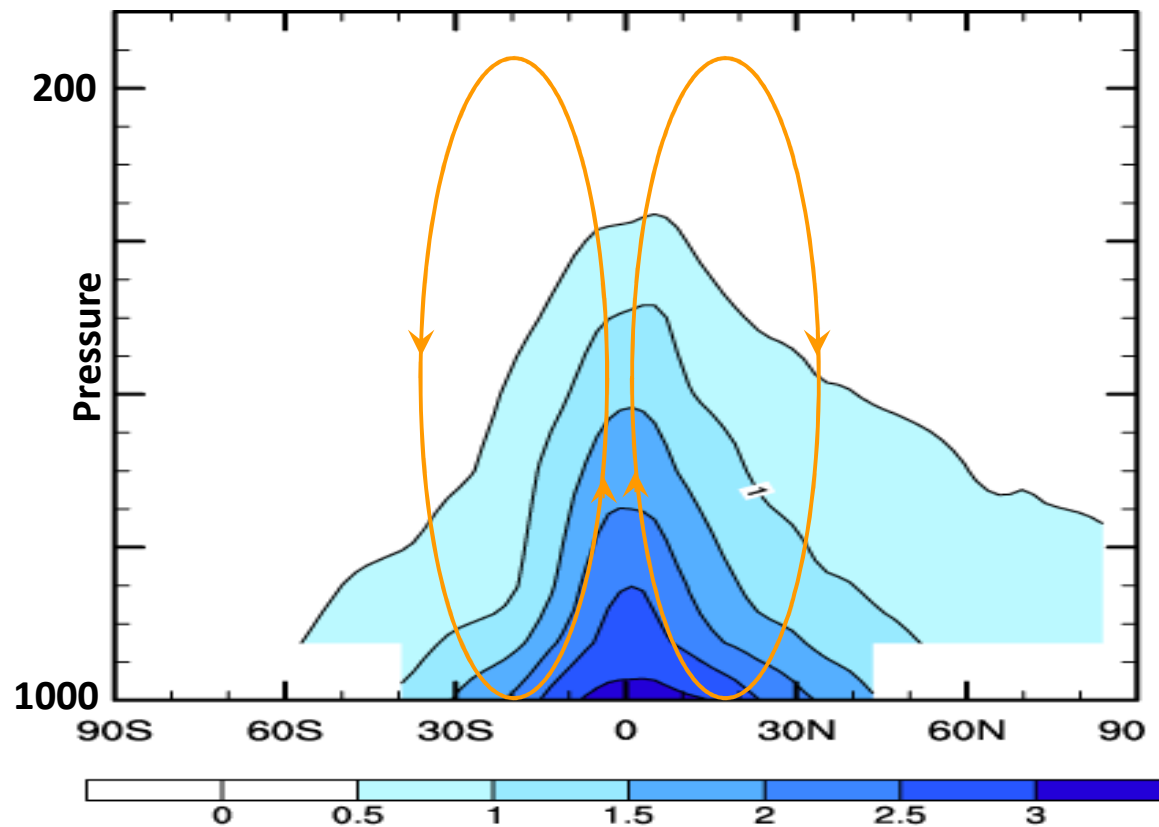


What determines rainfall change?

The wet gets wetter

(e.g., Neelin et al. 2003; Held & Soden 2006)

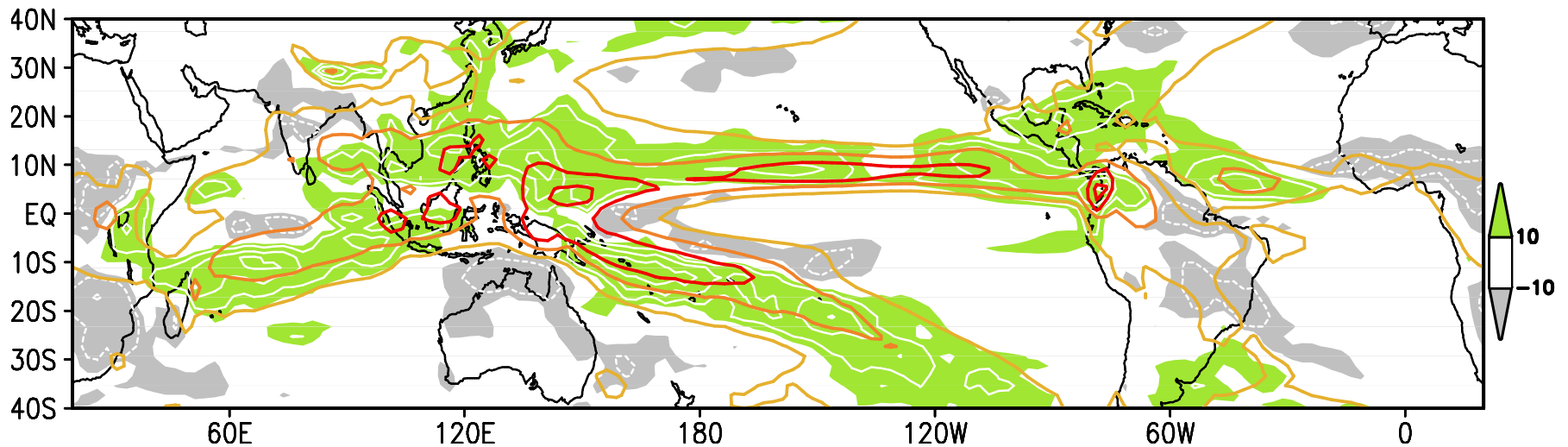
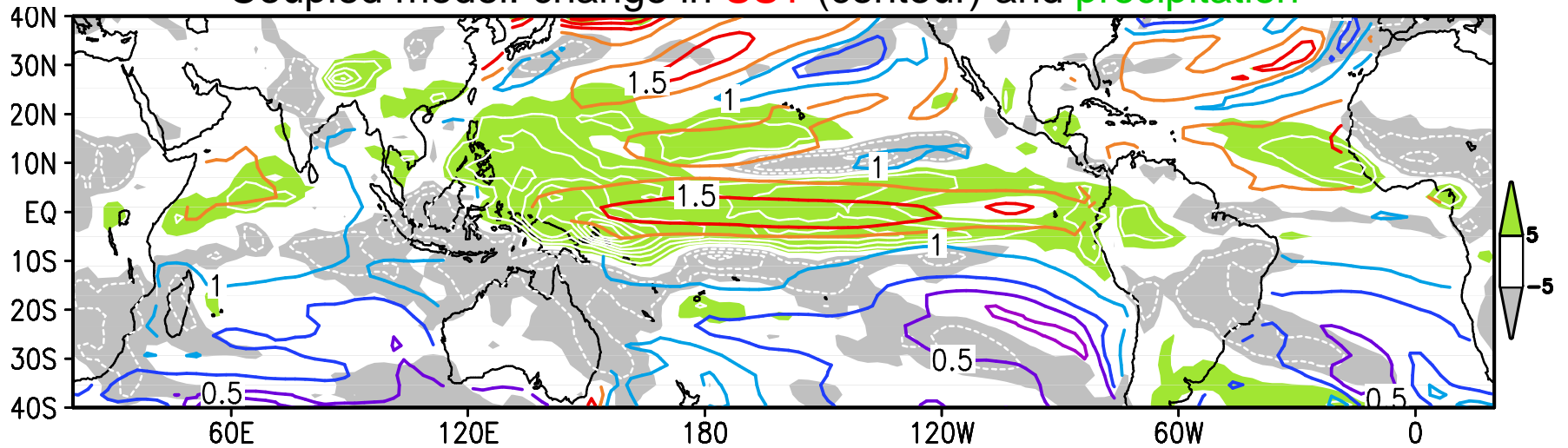
Precipitation increases in equatorial rain bands; decreases in subtropics; and increases in high-latitudes due to increase in moisture transport



Zonal-mean change in specific humidity δq [g/kg]

Warmer-get-wetter pattern

Coupled model: change in **SST** (contour) and **precipitation**

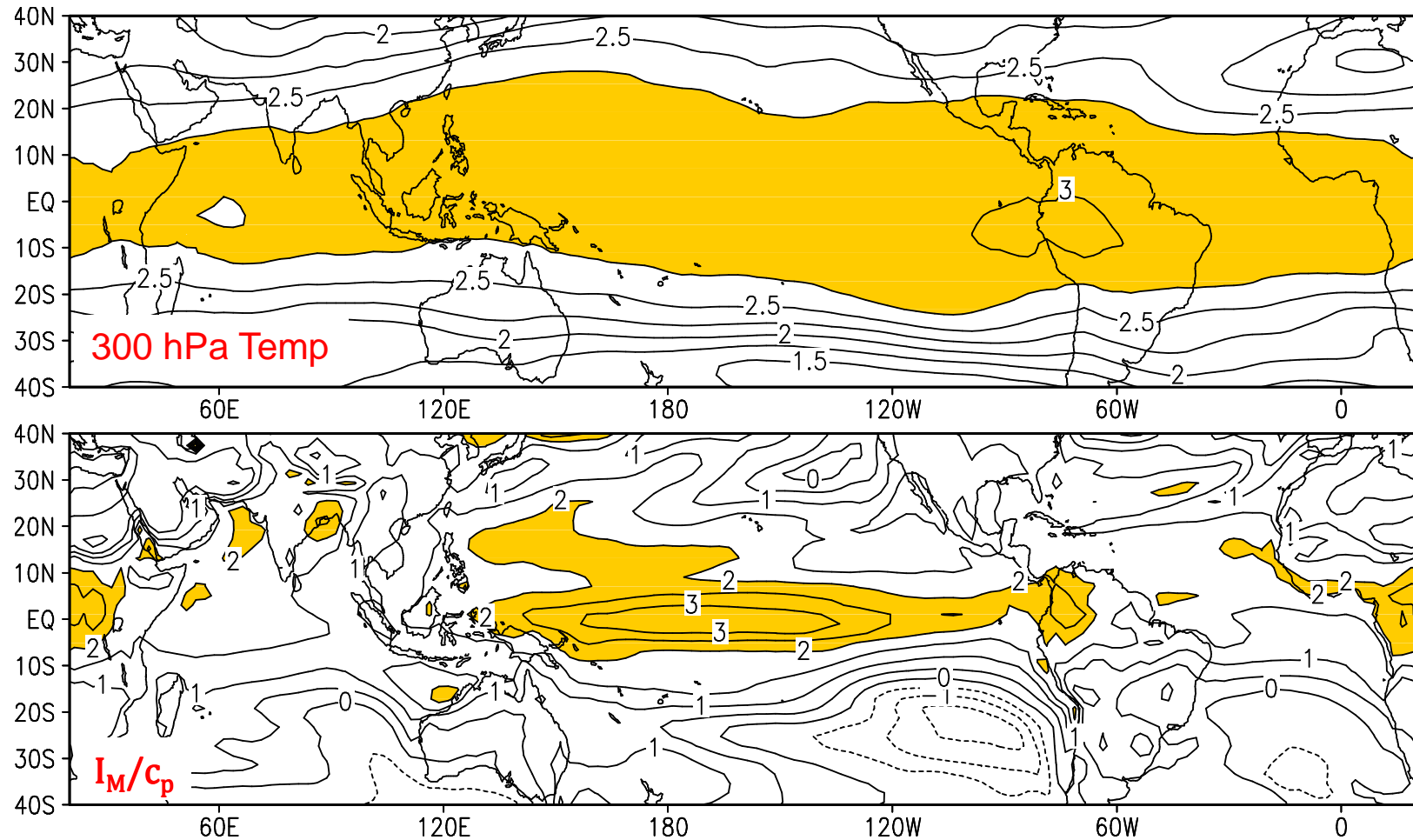


2K uniform SST warming: **mean** (contour) and **change** of precipitation

→ **Wet-get-wetter** pattern

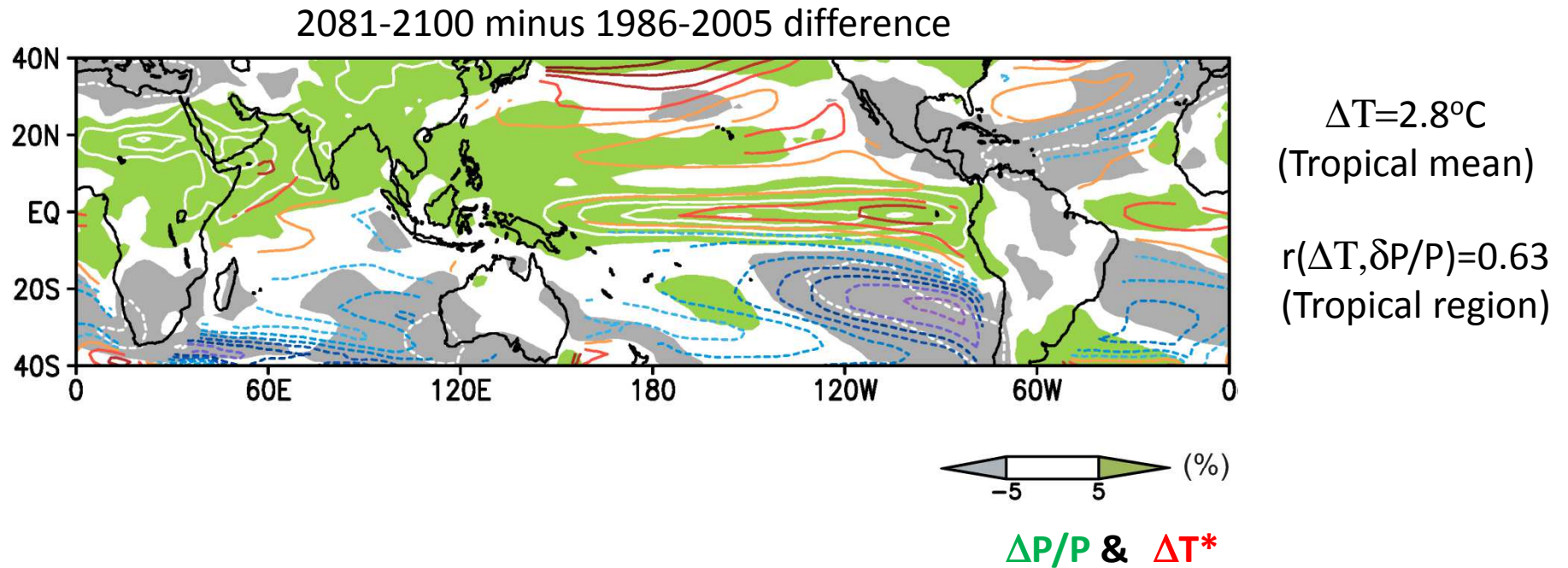
- Flat warming in upper troposphere ← weak temp gradient (Sobel et al. 2001)
- Convective instability follows closely SST patterns

Convective Instability: $I_M = (c_p T + Lq)_{\text{sfc}} - (c_p T + Lq)_{300 \text{ hPa}}$



Xie, S.-P., C. Deser, G.A. Vecchi, J. Ma, H. Teng, and A.T. Wittenberg, 2010: Global warming pattern formation: Sea surface temperature and rainfall. *J. Climate*, 23, 966-986.

Ocean warming pattern effect: **warmer get wetter**

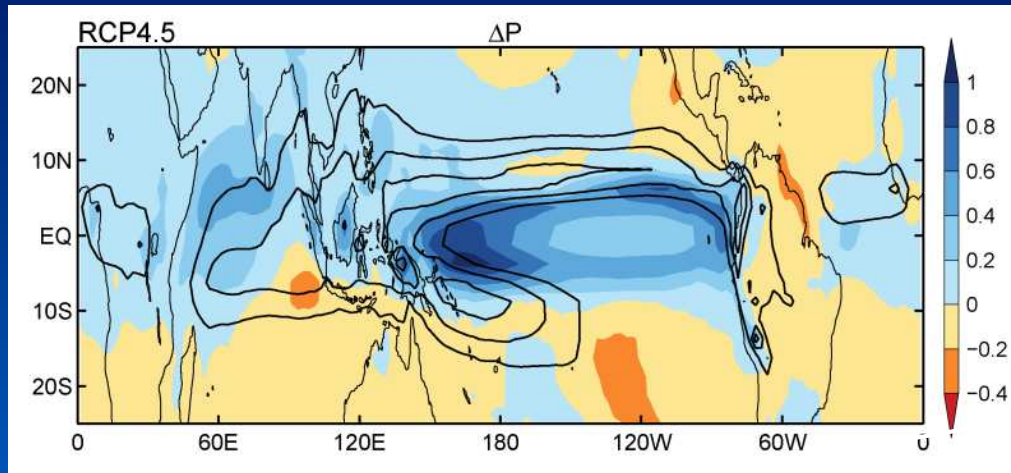


Robust patterns:

- Equatorial peak in Pacific and Atlantic
- Warmer NH than SH

IPCC AR5
Figure 14.8

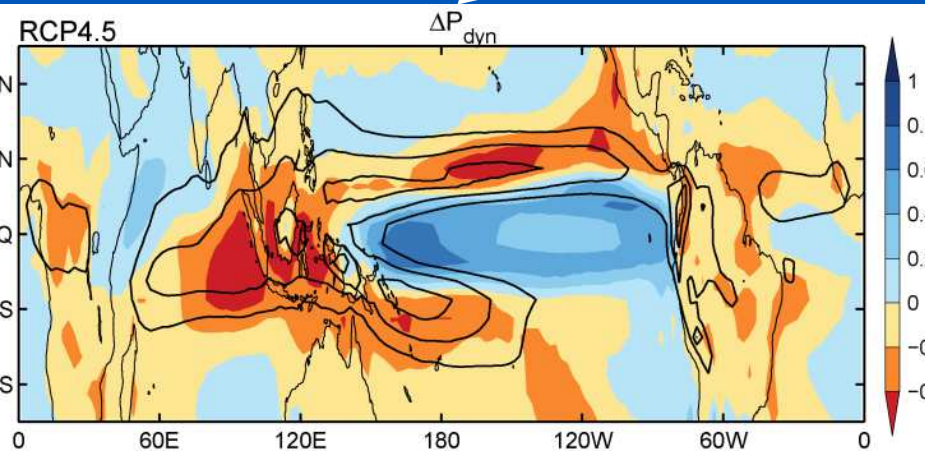
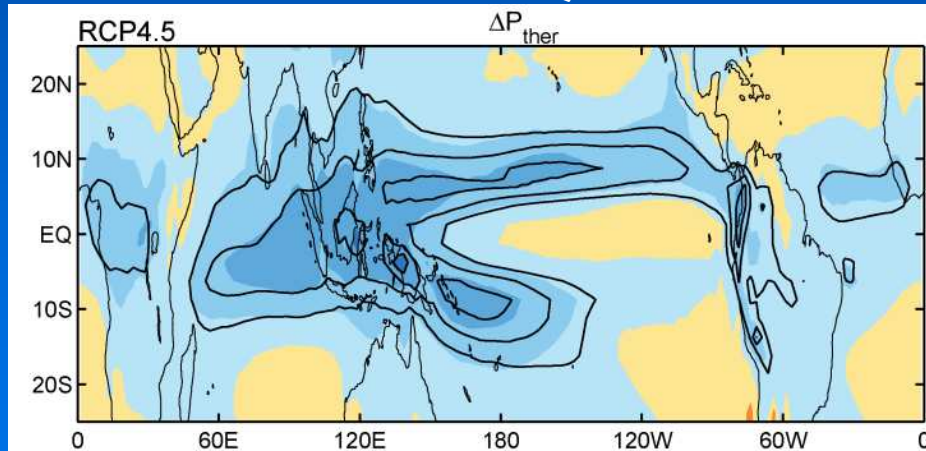
Decomposition of rainfall change



$$\Delta P \approx \Delta P_{ther} + \Delta P_{dyn}$$

$\omega \Delta q$

$\Delta \omega q$



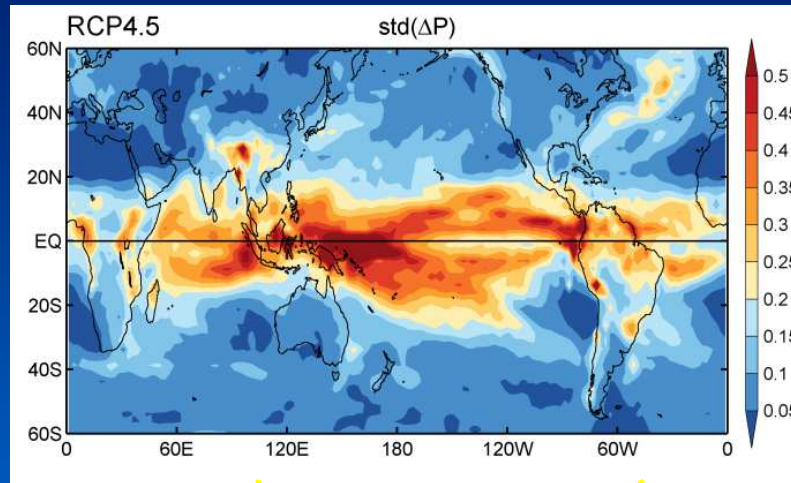
Wet get wetter

Circulation slowdown + SST pattern

Inter-model variability in 2050-99 minus 1950-99 difference

Dynamic component dominates inter-model spread

RCP4.5 Inter-model spread



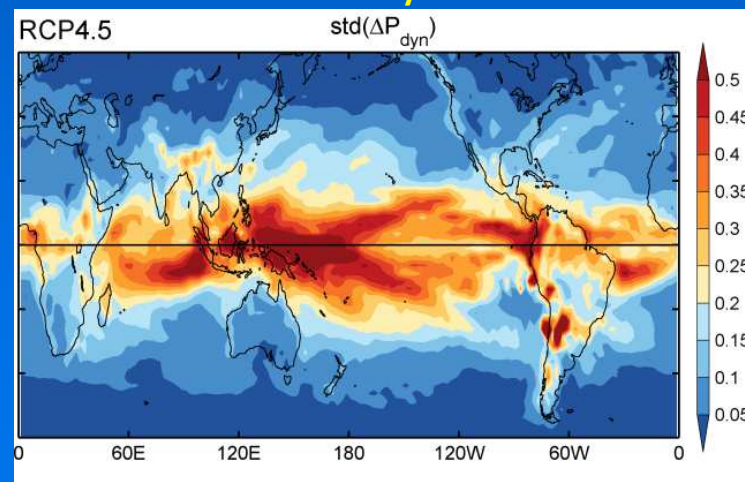
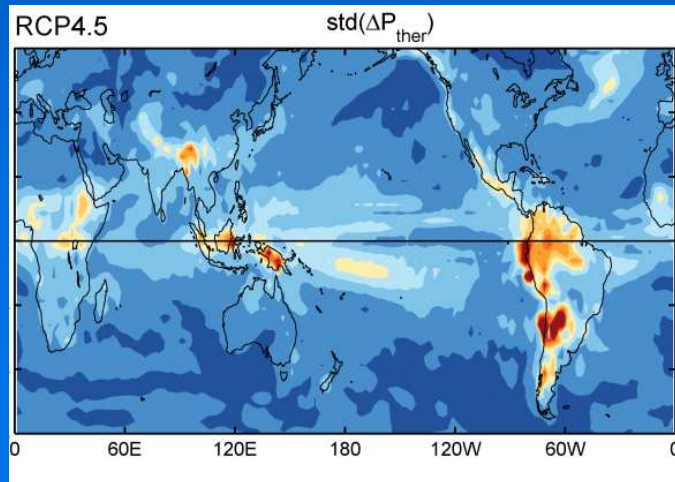
Circulation uncertainty
(air-sea coupling)

$$\Delta P \approx \underbrace{\Delta P_{ther}}_{\omega \Delta q} + \underbrace{\Delta P_{dyn}}_{\Delta \omega q}$$

ΔP_{ther}

ΔP

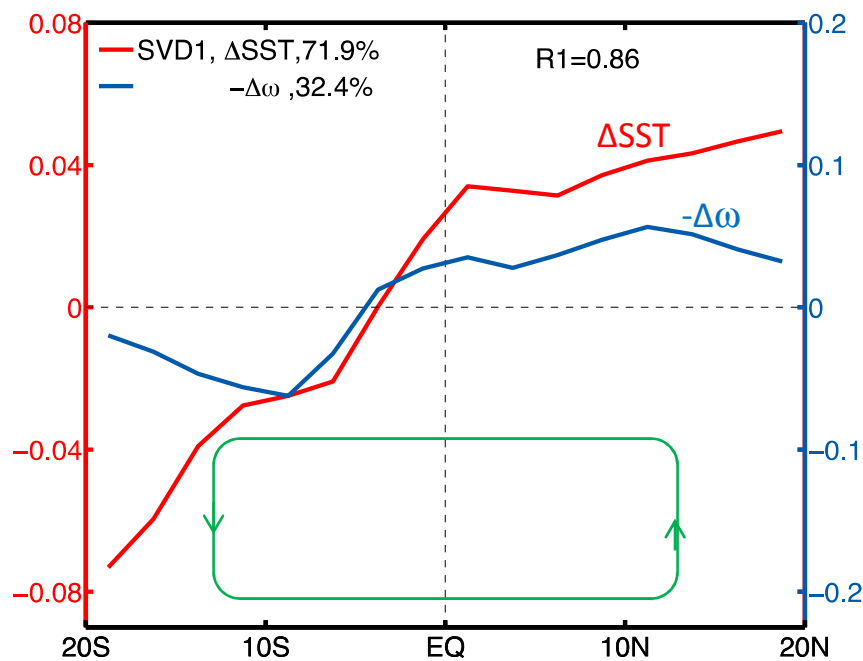
ΔP_{dyn}



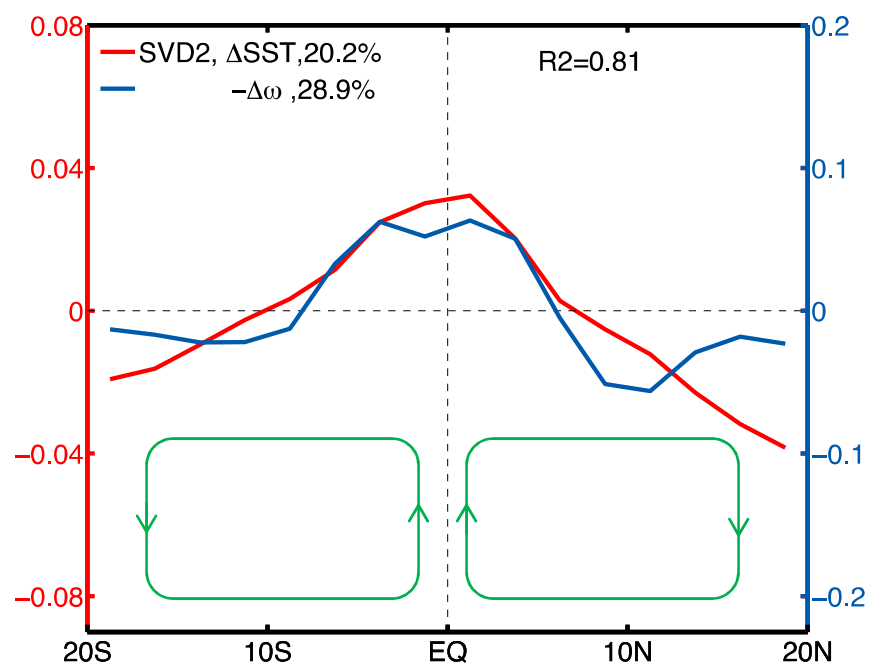
S.M. Long, in prep.

Atmospheric circulation uncertainty is tightly coupled with SST patterns

SVD analysis of inter-model spread of zonal mean Δ SST & $\Delta\omega$ in CMIP5



Uncertainty in cross-eq. gradient

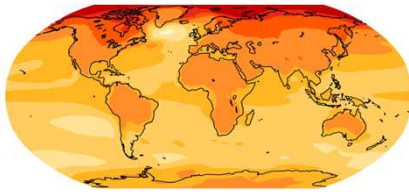


Uncertainty in eq. peak

See (Ma & Xie 2013, JC) for CMIP3 analysis

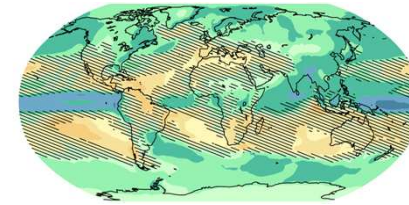
Leading uncertainties for **global mean** temperature

- Radiative forcing: aerosol
- Climate feedback: cloud



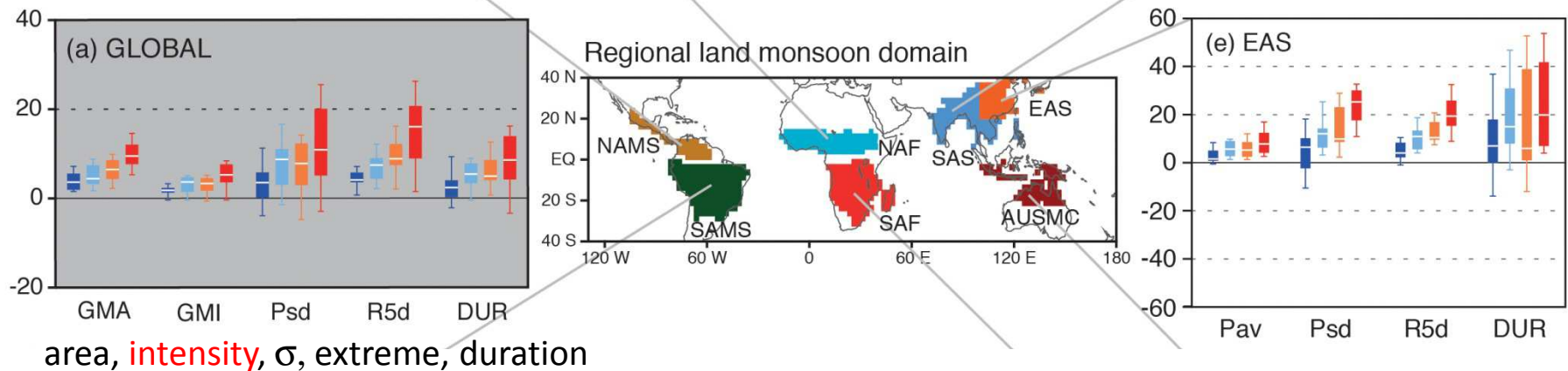
Leading uncertainty for **regional change**: atmospheric circulation

- Coupling with ocean in tropics
- Internal variability in mid-latitudes



Leading sources of global and regional climate projections. Insets are multi-model projections for surface temperature (left) and precipitation (right) from IPCC AR5.

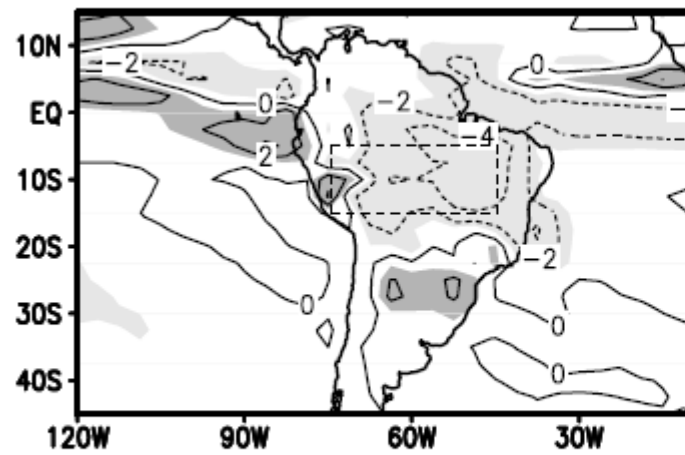
Monsoon change in AR5



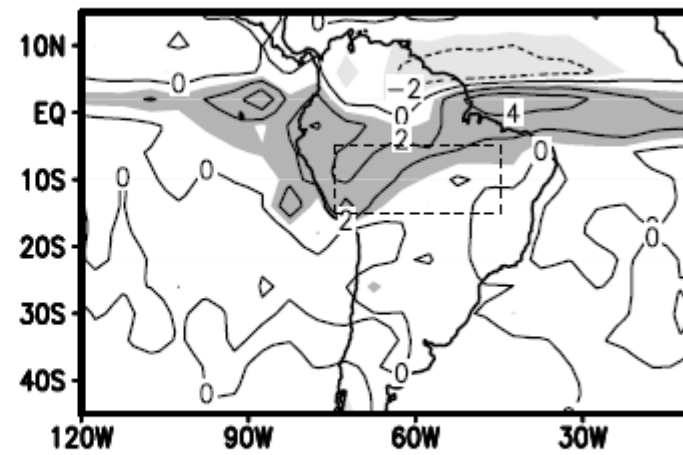
Globally, it is *likely* that the area encompassed by monsoon systems will increase over the 21st century. While monsoon winds are *likely* to weaken, **monsoon precipitation is *likely* to intensify due to the increase in atmospheric moisture.** ... resulting in lengthening of the monsoon season in many regions. AR5 SPM

Fig. TS-24

c) HadCM3 Dec

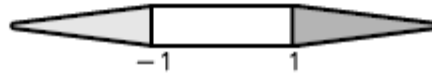


d) GISS-ER Dec

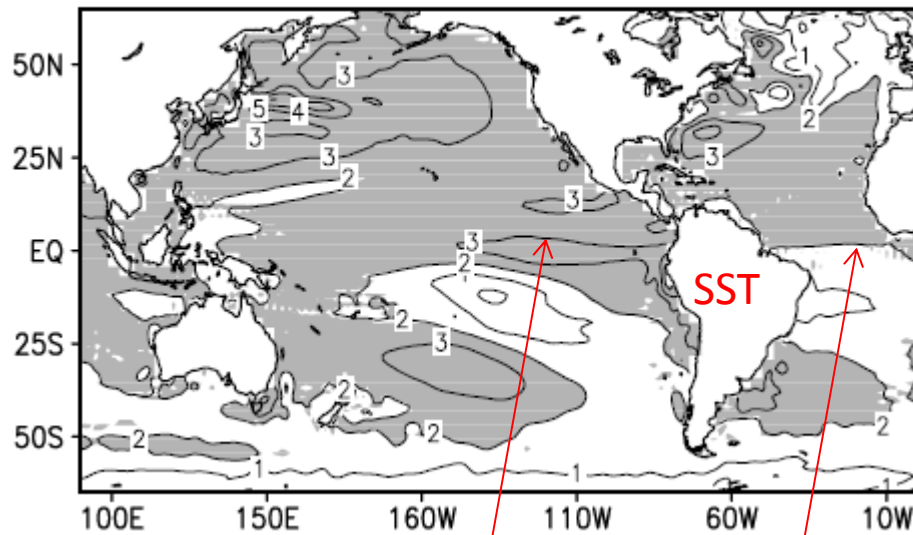


South American monsoon

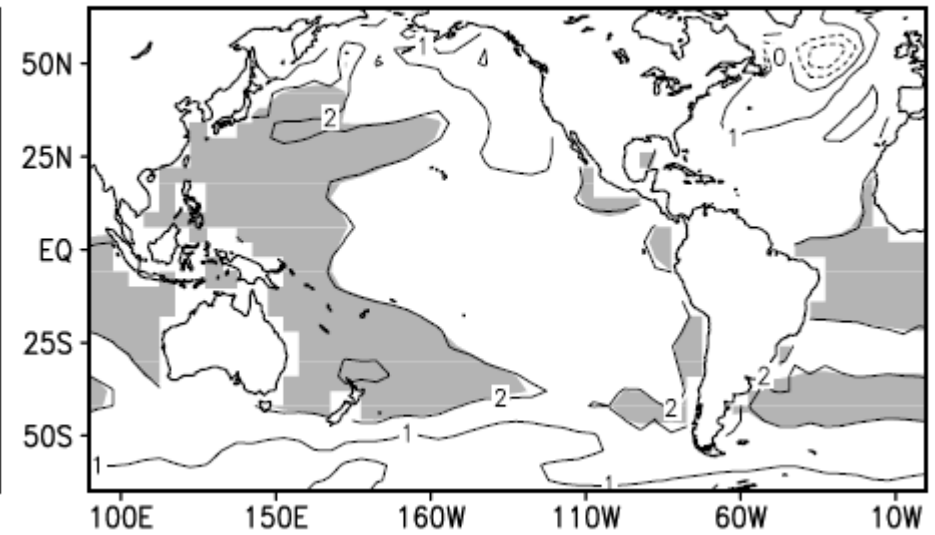
Rainfall change



a) HadCM3



b) GISS-ER



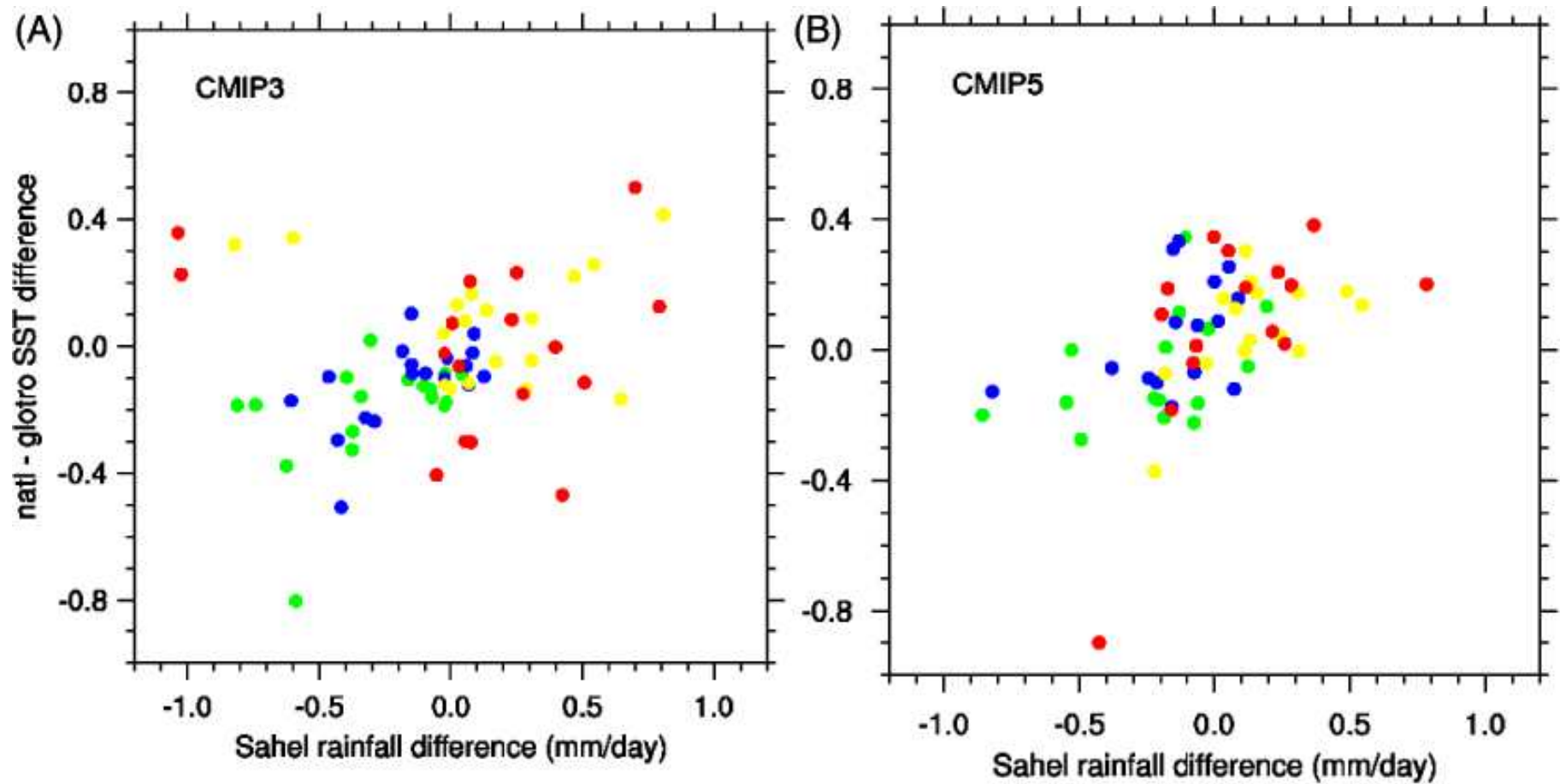
El Niño-like

NS SST grad

W. Li et al. 2006 (JGR)

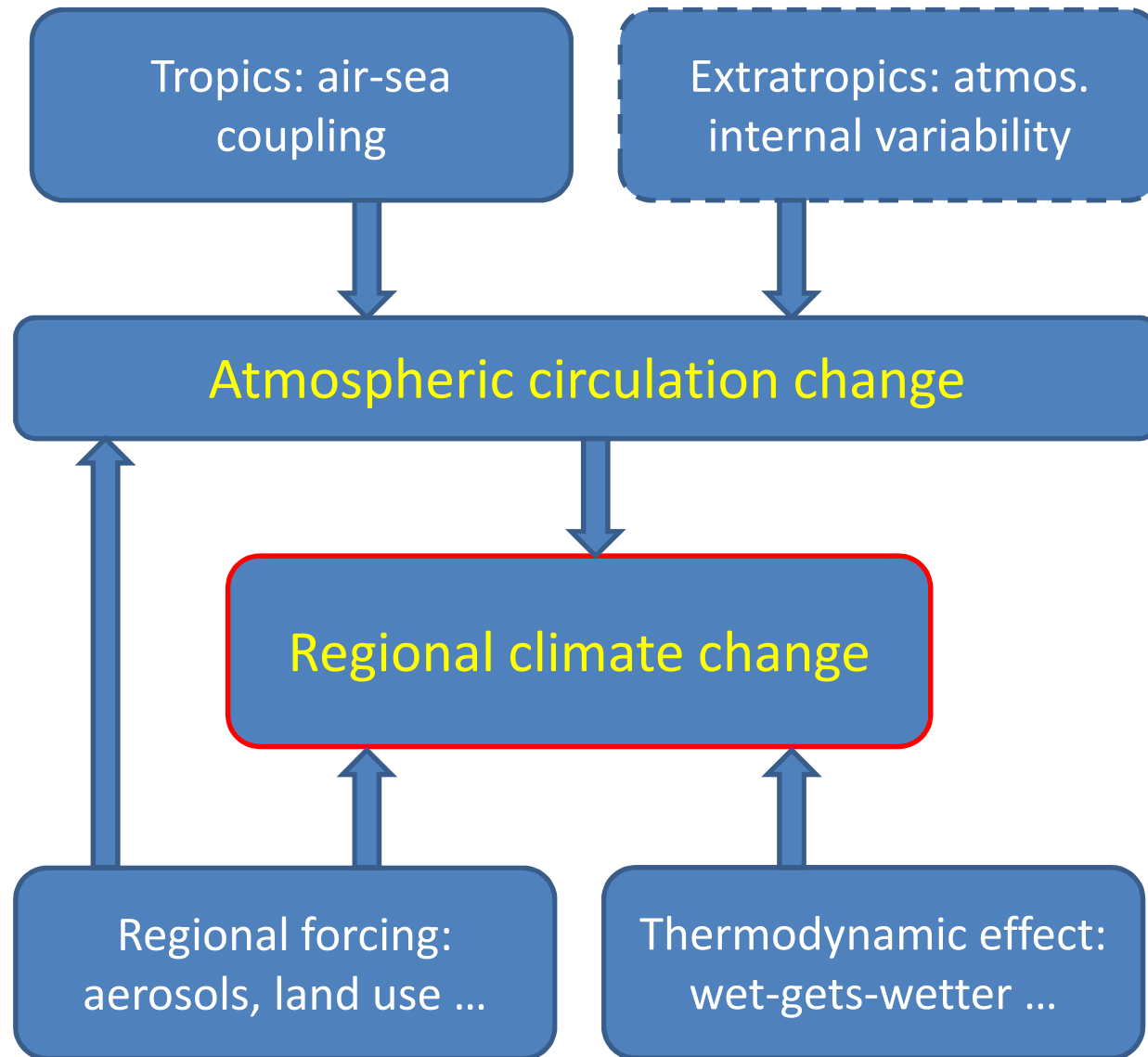
West African monsoon

Sahel rainfall change vs. N Atlantic relative SST



Giannini et al. (2013, ERL)

Physical origins of regional climate change.



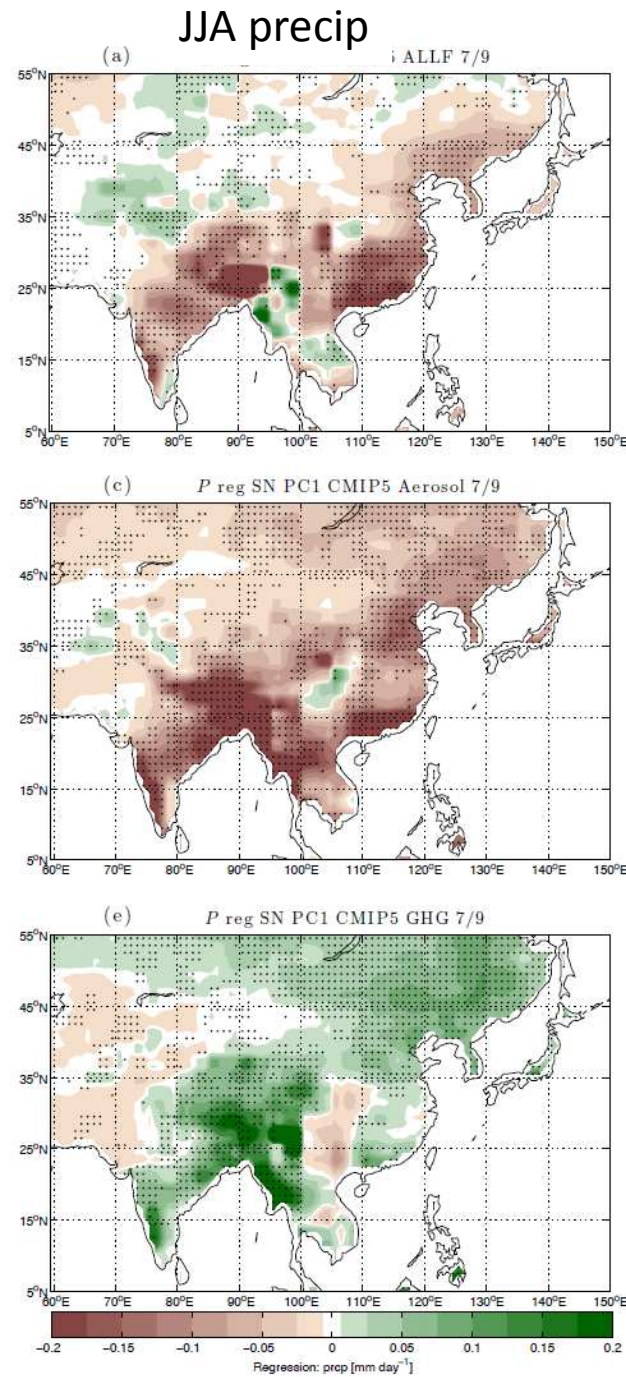
Near emission source,
aerosol forcing
dominates circulation
and rainfall response.

All forcing

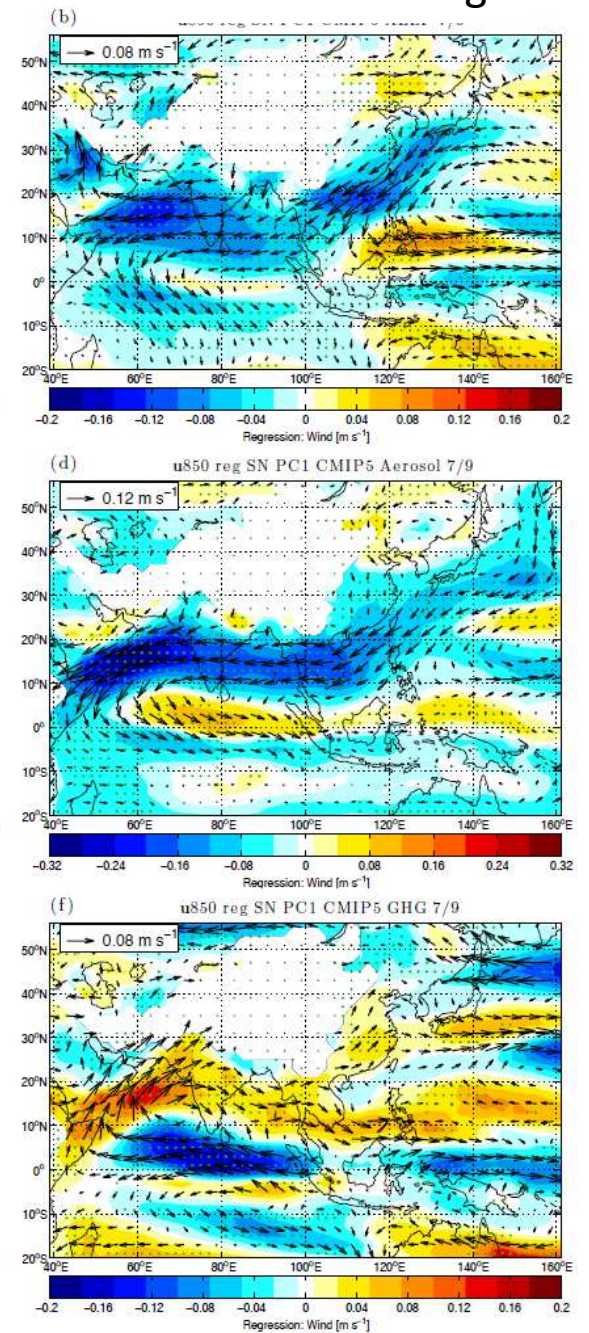
Asian summer monsoon

Aerosol

GHG



850 hPa wind change



Water budget

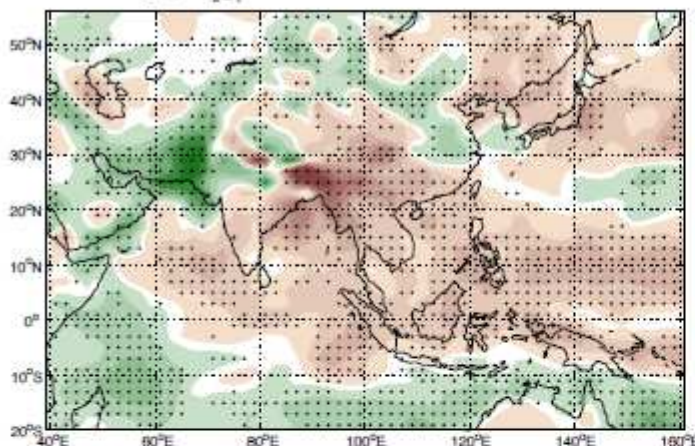
Aerosol run: Large dynamic component

GHG run: Large thermodynamic component

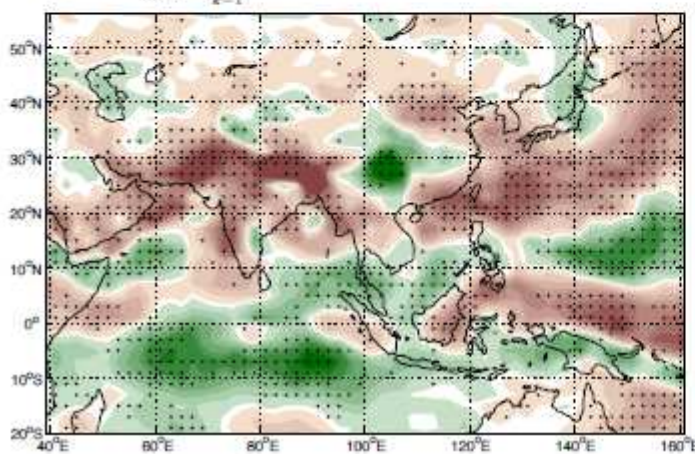
thermodynamic component

dynamic component

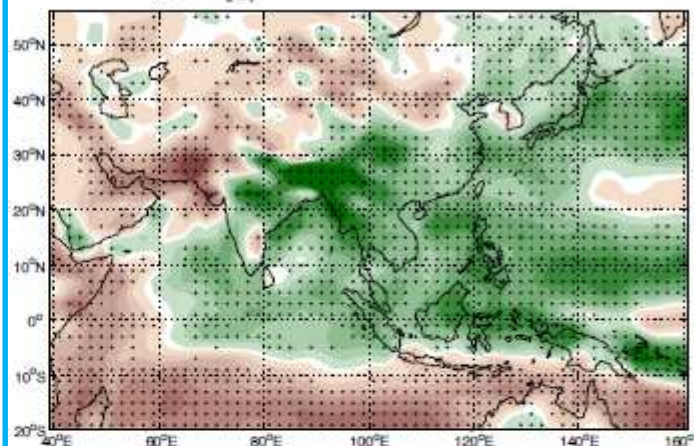
$$(e) \quad -\frac{1}{g\rho_w} \nabla \cdot \sum_{k=1}^{10} \bar{\mathbf{u}}_k \bar{\mathbf{q}}_k \Delta \bar{p}_k \text{ reg SN PC1 CMIP5 Aerosol 7/9}$$



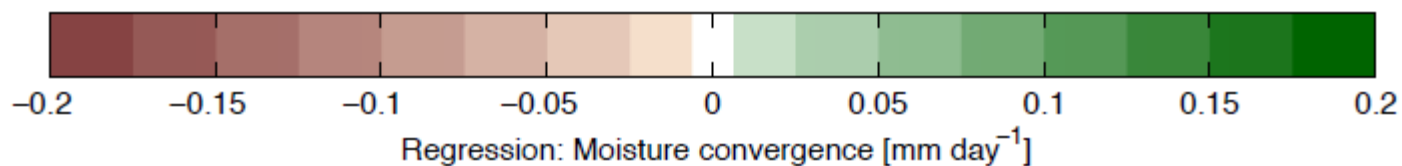
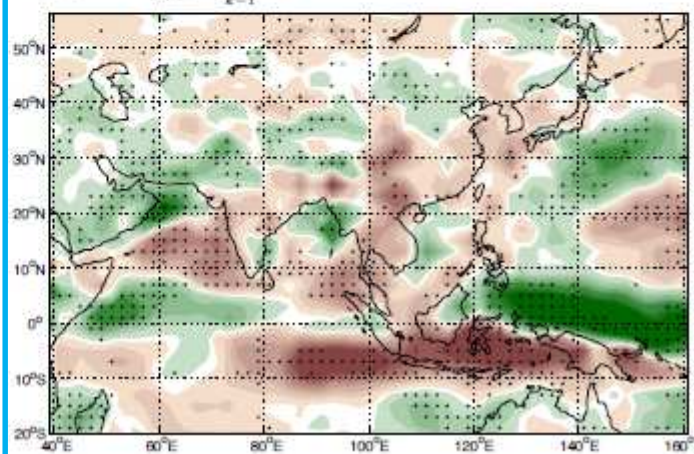
$$(h) \quad -\frac{1}{g\rho_w} \nabla \cdot \sum_{k=1}^{10} \bar{\mathbf{u}}_k \bar{\mathbf{q}}_k \Delta \bar{p}_k \text{ reg SN PC1 CMIP5 Aerosol 7/9}$$



$$(f) \quad -\frac{1}{g\rho_w} \nabla \cdot \sum_{k=1}^{10} \bar{\mathbf{u}}_k \bar{\mathbf{q}}_k \Delta \bar{p}_k \text{ reg SN PC1 CMIP5 GHG 7/9}$$



$$(i) \quad -\frac{1}{g\rho_w} \nabla \cdot \sum_{k=1}^{10} \bar{\mathbf{u}}_k \bar{\mathbf{q}}_k \Delta \bar{p}_k \text{ reg SN PC1 CMIP5 GHG 7/9}$$



Li, Ting et al.
(2015, JC)

Challenges

- Over ocean, SST pattern, circulation and rainfall are coupled (diagnostic), but are there overarching principals (predictive)?
- What determines rainfall/monsoon change over continents?
- Interannual variability and extremes?