

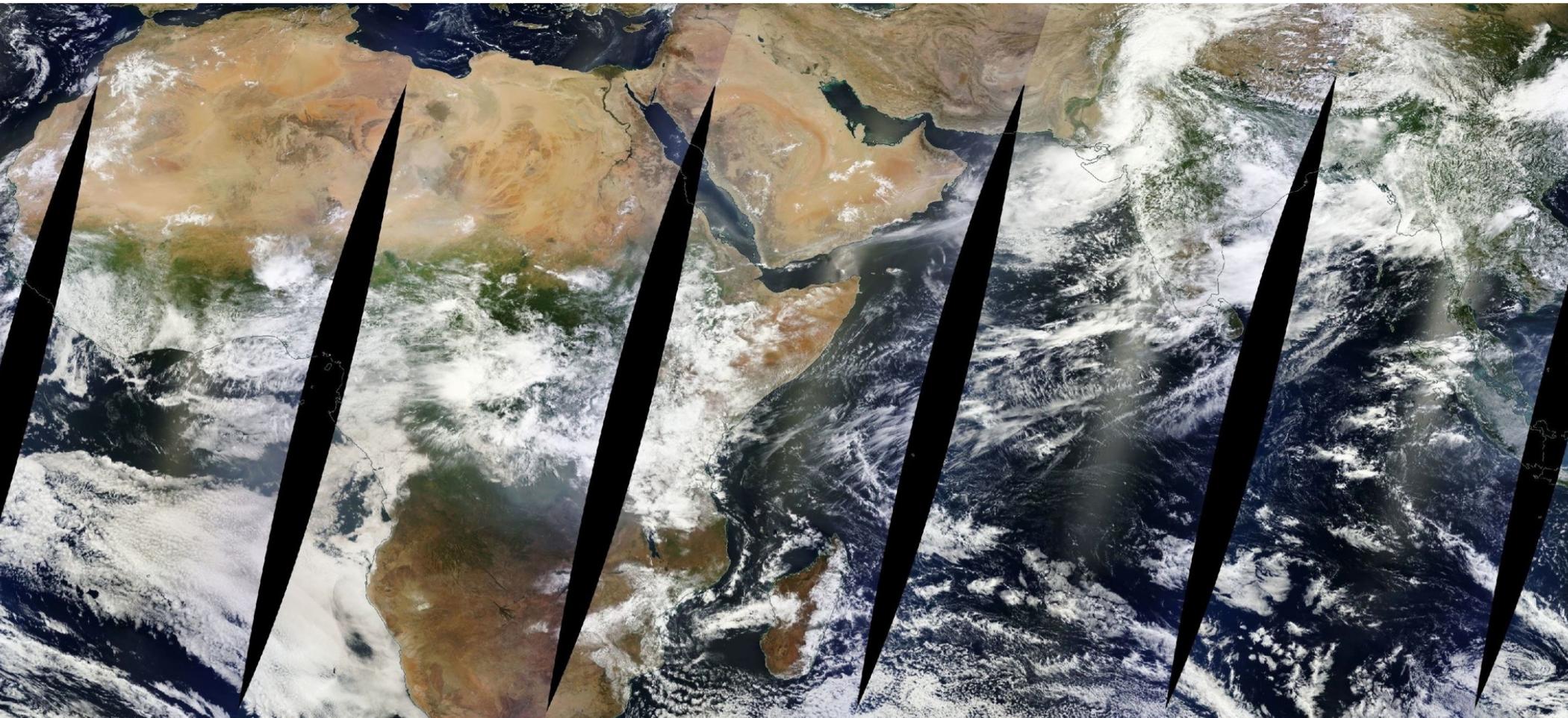
A review of energetic and dynamic theories of monsoons

William Boos
May 18, 2015

Yale

contributions from: John Hurley,
Ravi Shekhar, Trude Storelvmo

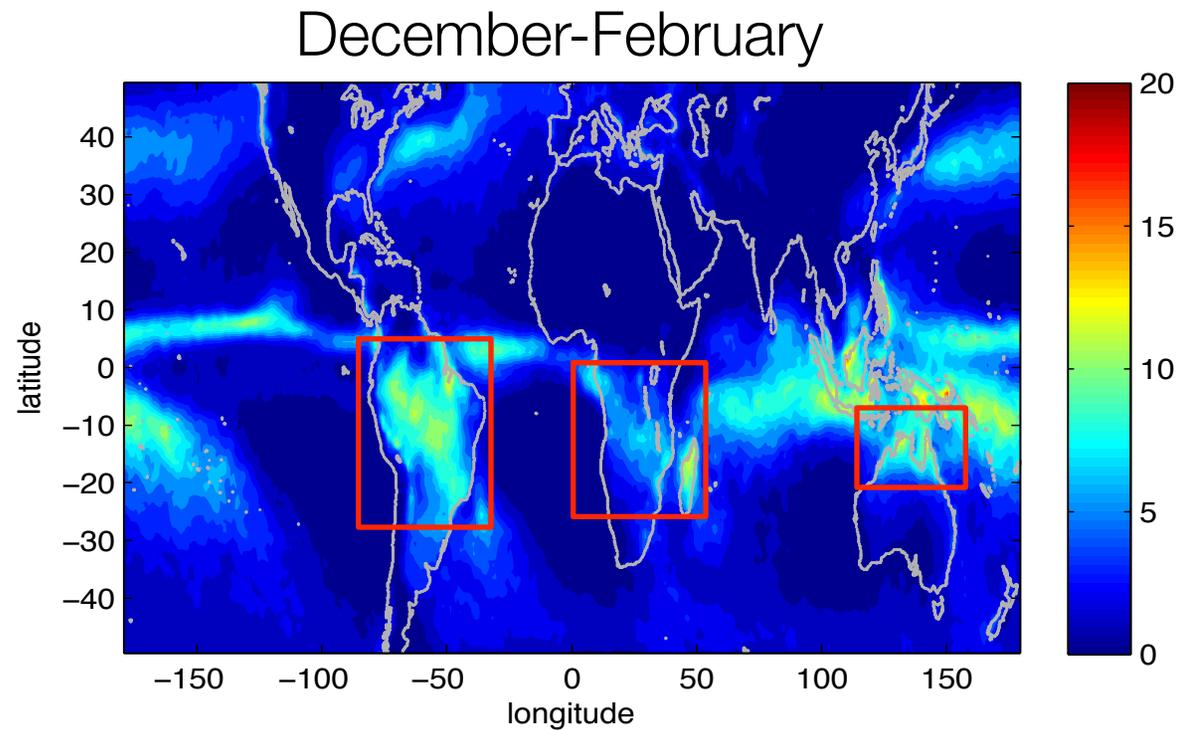
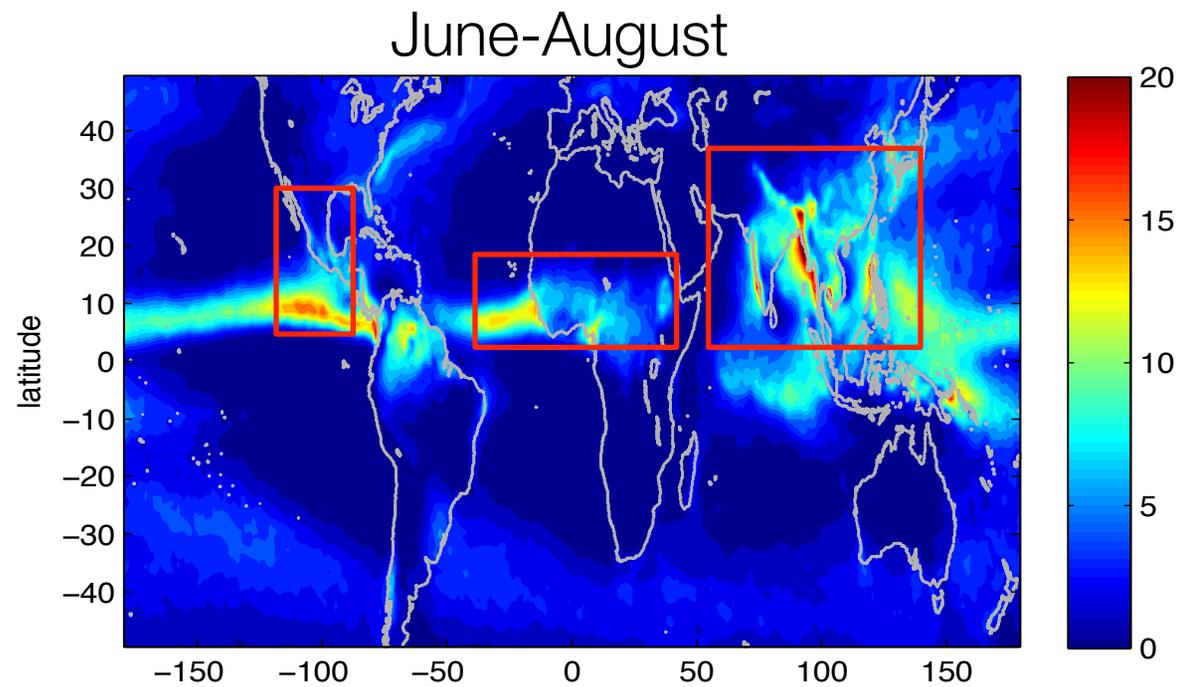
Financial
support:



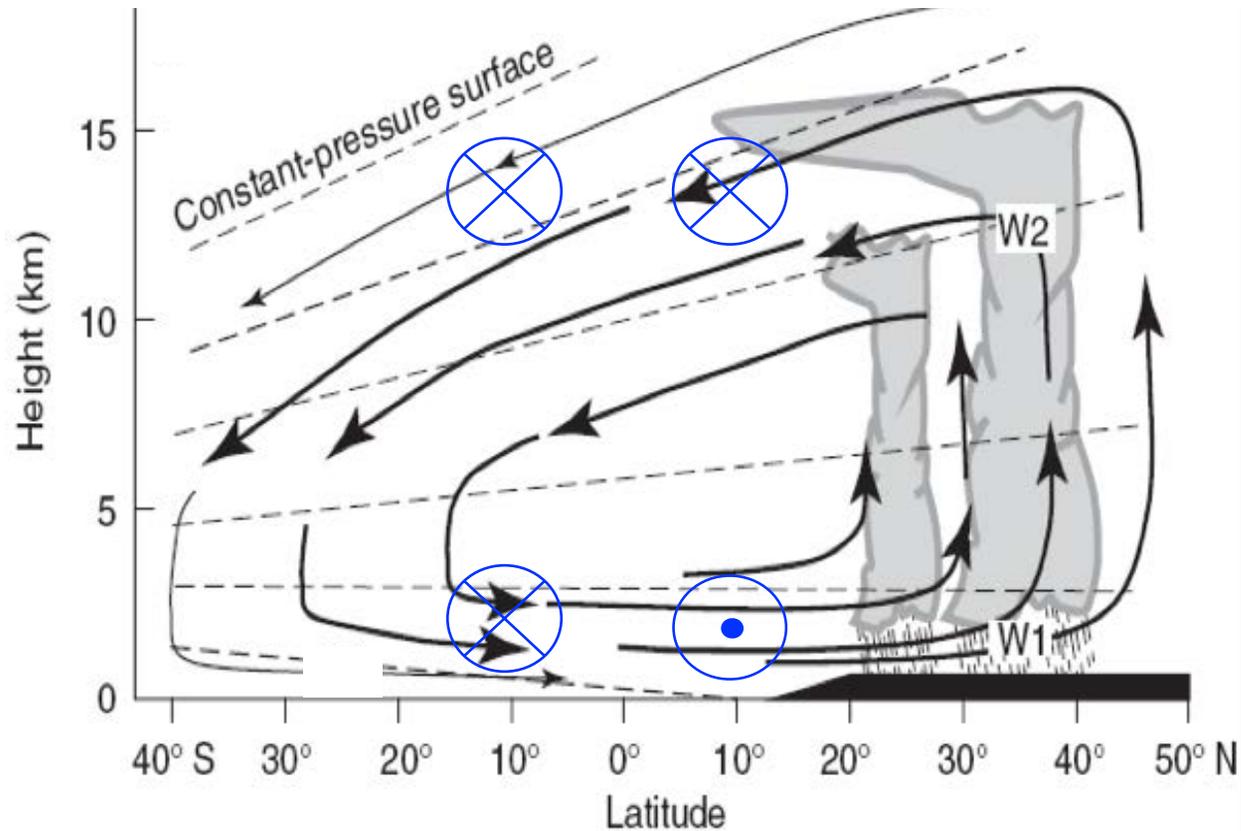
Sept. 4, 2014 Terra/MODIS true color reflectance from NASA's EOSDIS Worldview

The global monsoon

Globally, a large fraction of rain falls in monsoon regions during solstice seasons



An idealized view of monsoons

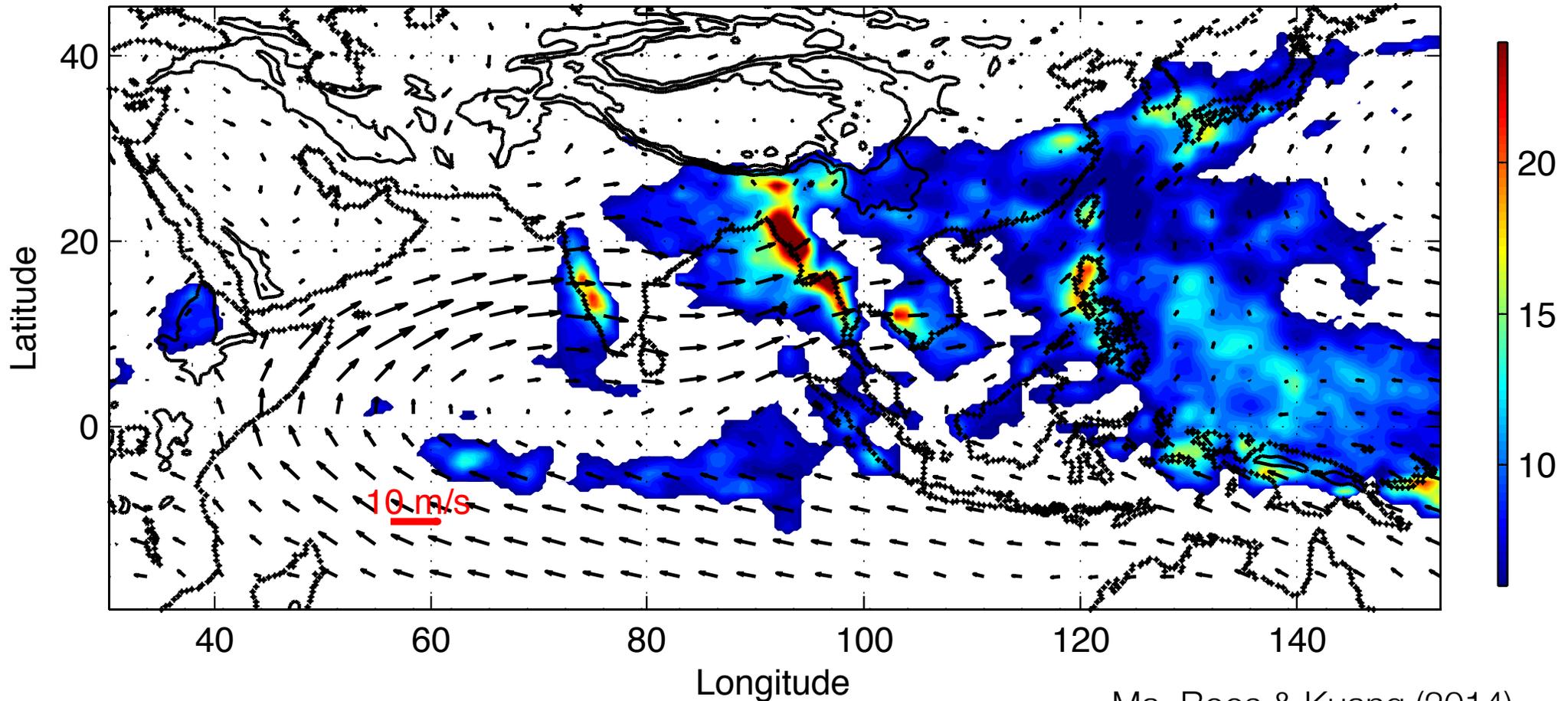


Webster and Fasullo (2003)

- precipitation occurs in ascending air in summer hemisphere
- circulation is closed by cross-equatorial flow and descent in winter hemisphere
- reversal of typical “trade winds” in summer hemisphere

This 2D view hides much complexity, but accurately describes zonal mean structure

JJA climatology of TRMM precipitation (mm/day)
and ERA-40 850 hPa wind



Ma, Boos & Kuang (2014)

Going beyond the descriptive: how can we ***understand*** the variability of monsoons?

Limit our focus to seasonal-mean behavior:

- Interannual variability (2002 drought in India)
- Decadal variability (1970s-80s Sahel drought; past 30-year drying trend in East Africa)
- Response to anthropogenic forcings (greenhouse gases, aerosols, land use change)
- Paleoclimate variability (mid-Holocene North African humid period)

How to think about monsoon circulations?

i.e. given a forcing, how will the precipitating circulation respond?

Dominant categories of theory:

- Surface temperature gradients drive flow **wrong**

- Convective quasi-equilibrium **mechanistic but diagnostic**

- Vertically integrated atmospheric energy budget **prognostic but not mechanistic**

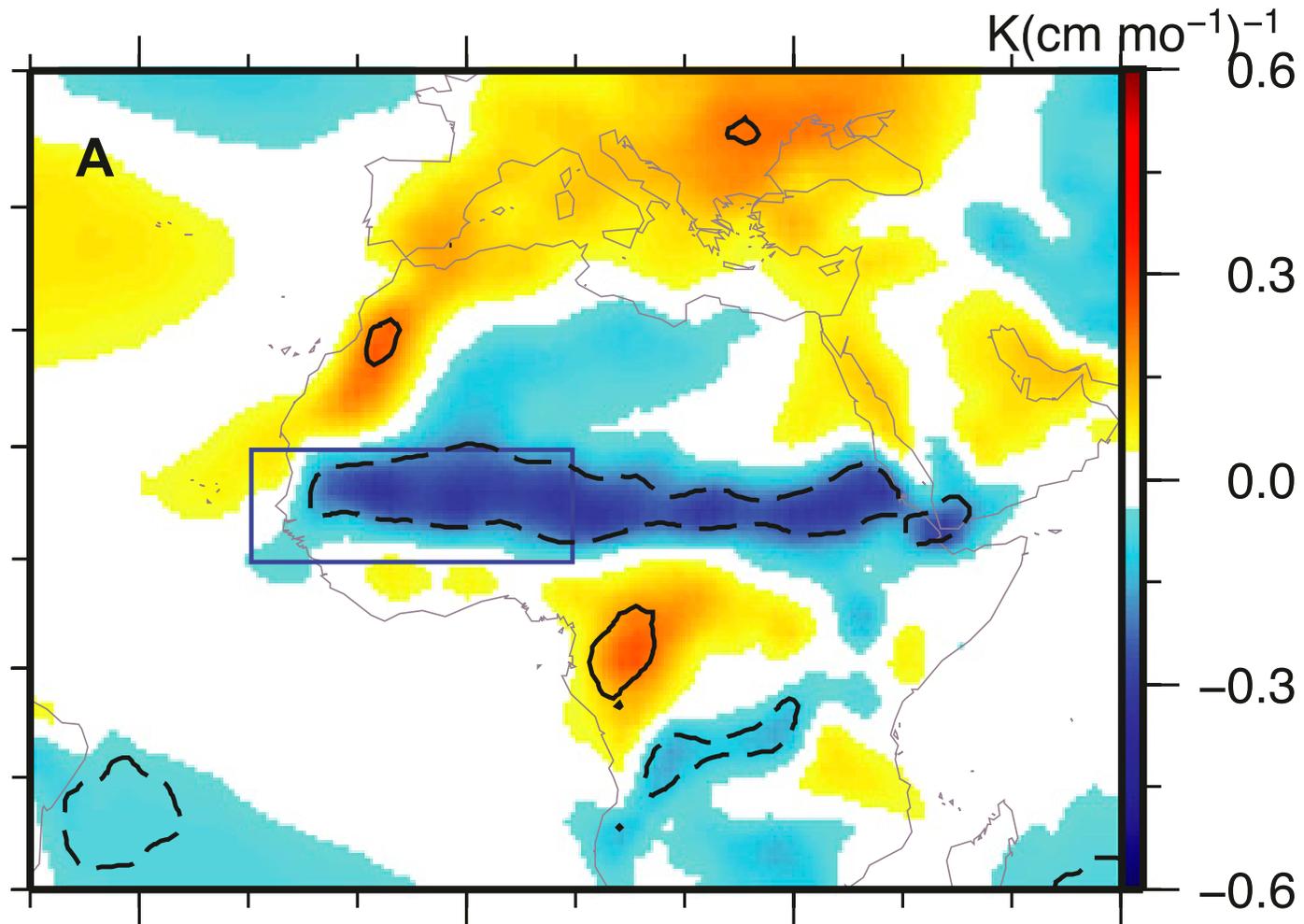
Disproof of idea that surface temperature gradients drive flow

regression of near-surface air temperature on Sahel rainfall

Data: GHCN rainfall & ERA-Interim near-surface air temperature

Black contours surround regions significant at 5% level

Hurley & Boos (2013)



How to think about monsoon circulations?

i.e. given a forcing, how will the precipitating circulation respond?

Dominant categories of theory:

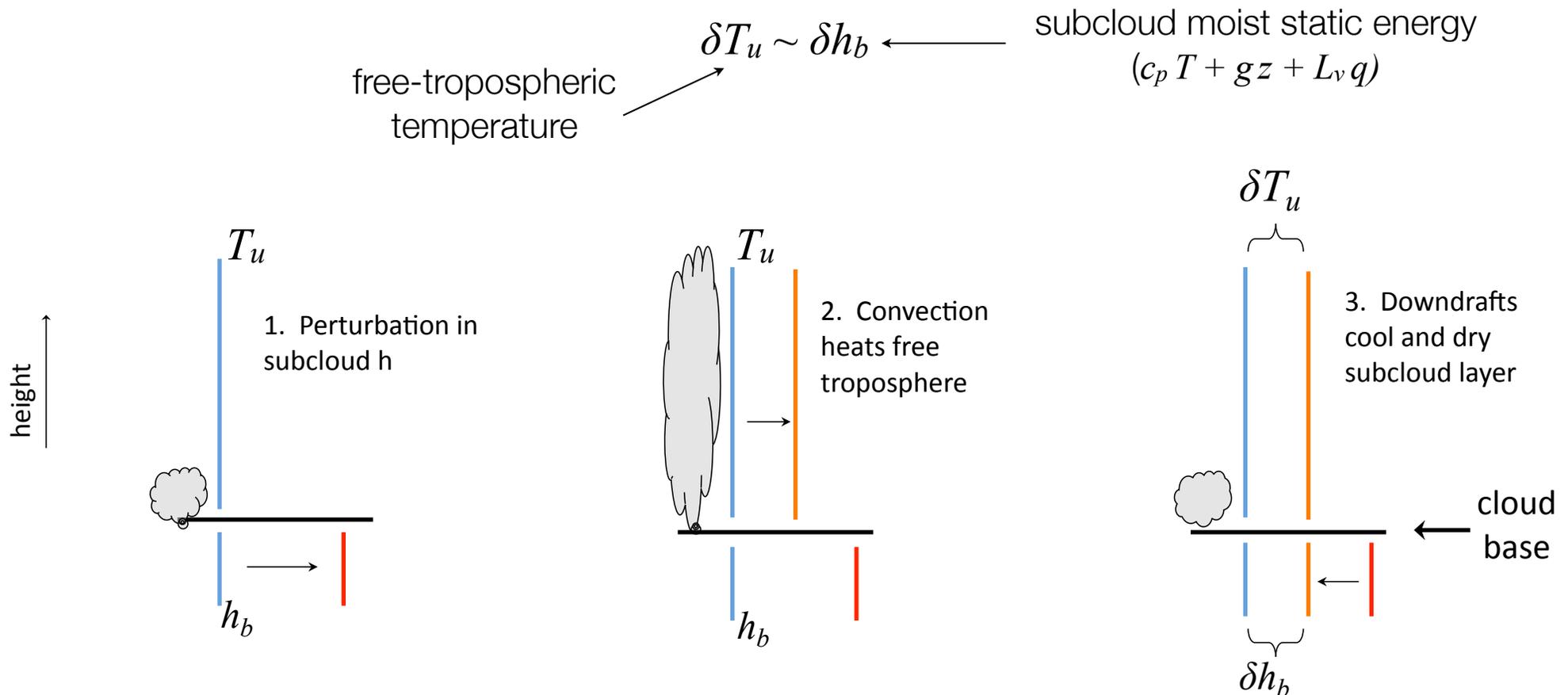
- Surface temperature gradients drive flow **wrong**

- Convective quasi-equilibrium **mechanistic but diagnostic**

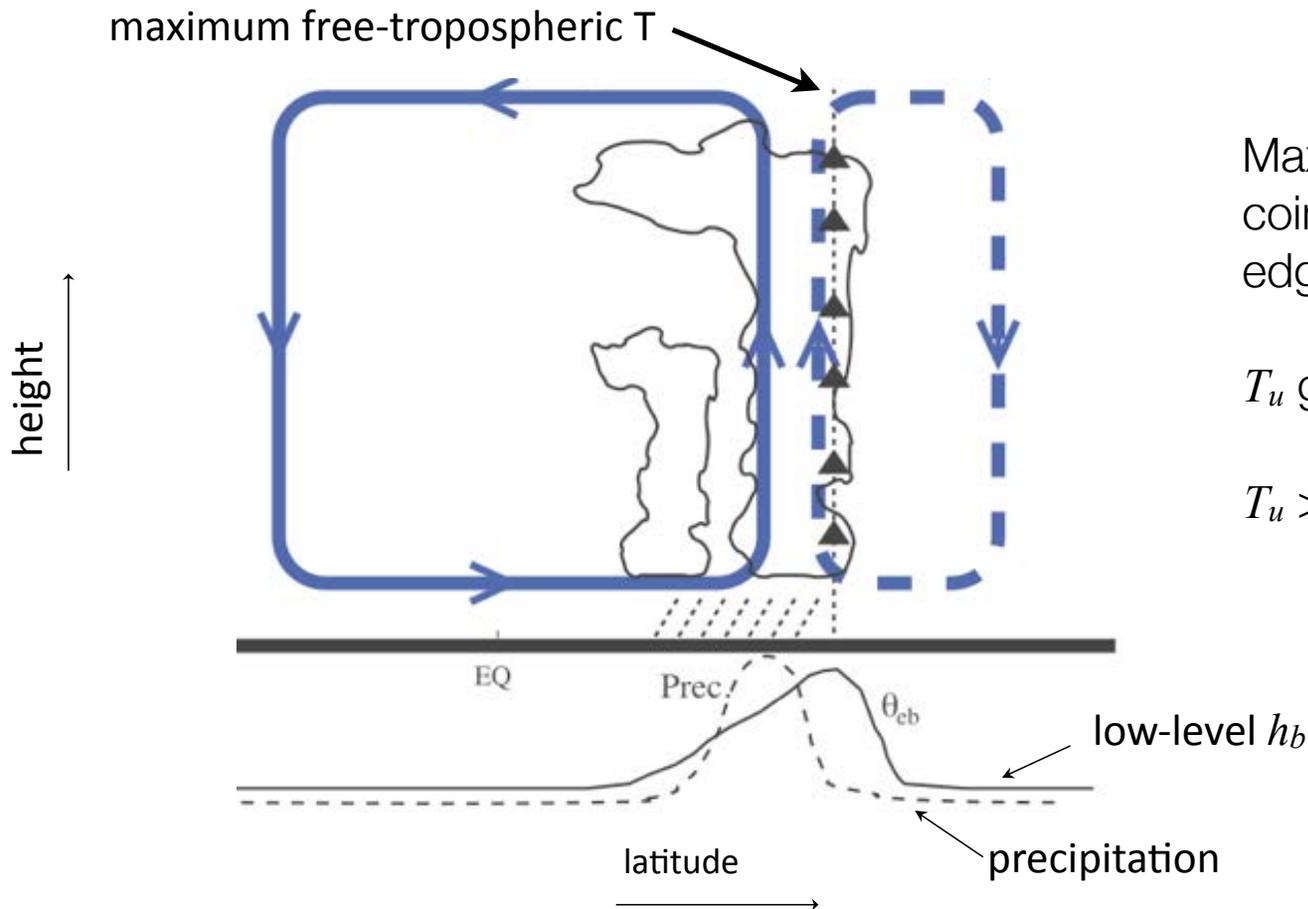
- Vertically integrated atmospheric energy budget **prognostic but not mechanistic**

Convective quasi-equilibrium: a modern way to think about convectively coupled large-scale flow

- Moist convection does not act as “heat source” for large-scale flow, but maintains free-troposphere near moist adiabat
- Changes in free-tropospheric temperature are in equilibrium with changes in subcloud entropy:



Convective quasi-equilibrium monsoons



Maxima of h_b and T_u should be coincident in monsoons, at poleward edge of cross-equatorial cell.

T_u gradients are weak

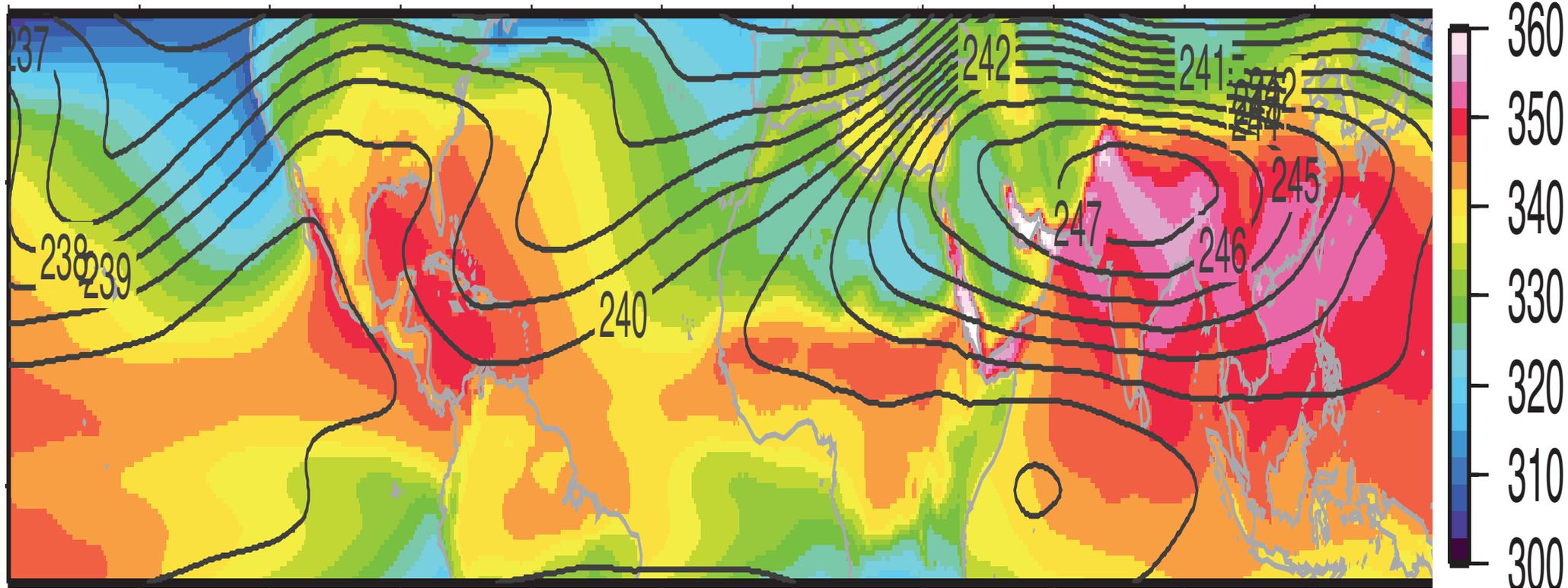
$T_u > h_b$ away from the ascent branch

Convective quasi-equilibrium framework describes South Asian summer mean state

July climatology

Contours: 200-400 hPa temperature (K)

Colors: surface air moist static energy ($c_p T + gz + L_v q$), in K

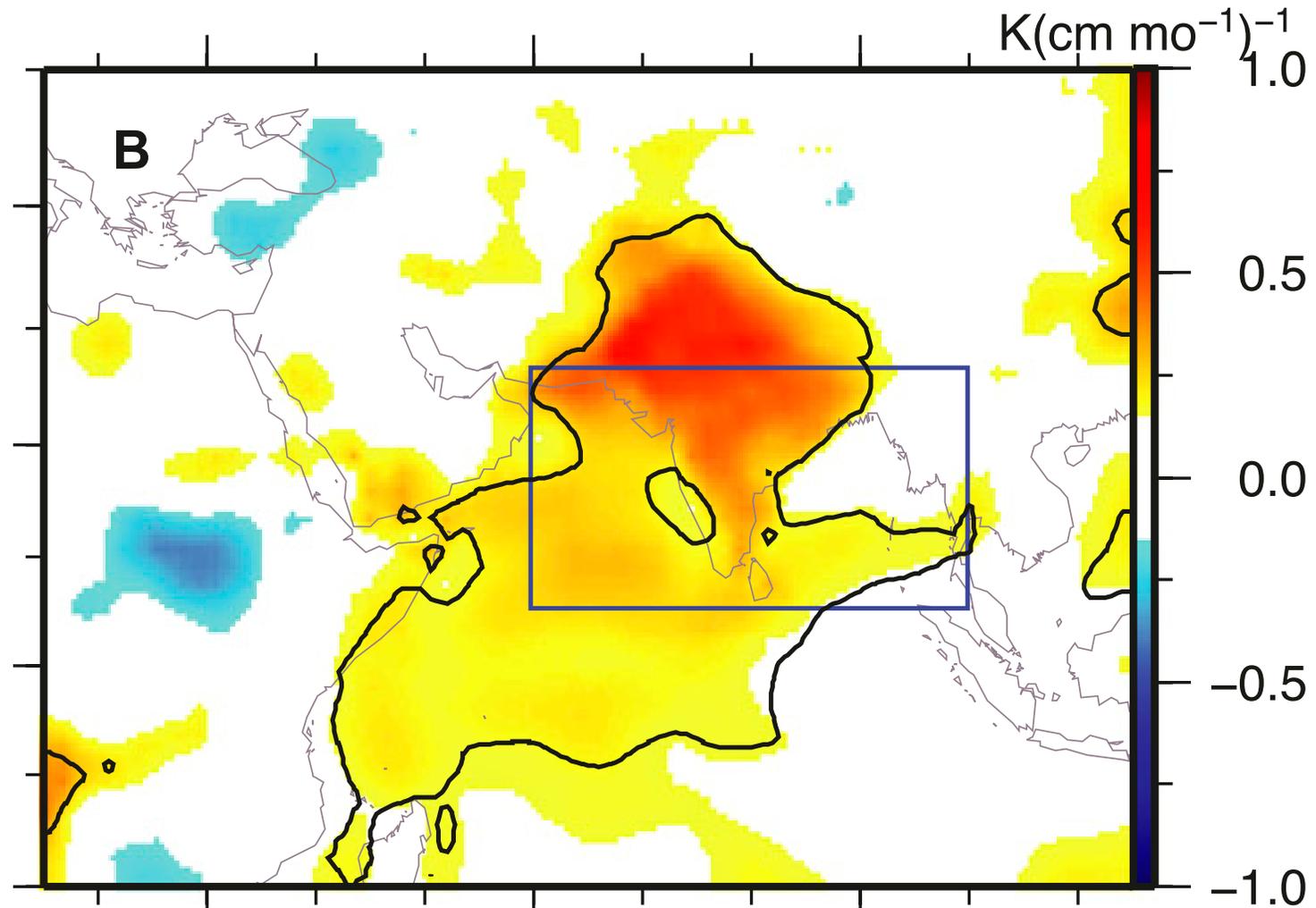


Data: ERA-Interim

Boos & Hurley (2013)

Convective quasi-equilibrium framework is also consistent with South Asian interannual variability

regression of near-surface moist static energy on monsoon rainfall

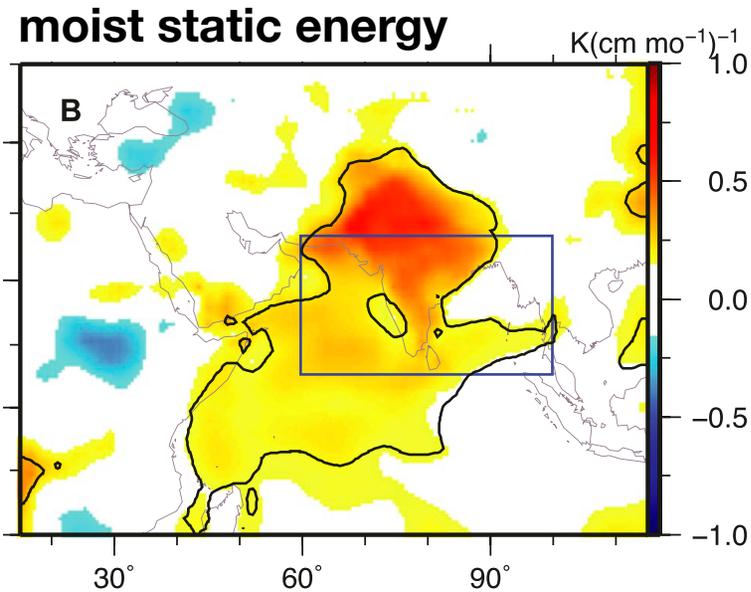


Data: GHCN
rainfall & ERA-
Interim near-
surface
equivalent
potential
temperature

Hurley & Boos
(2013)

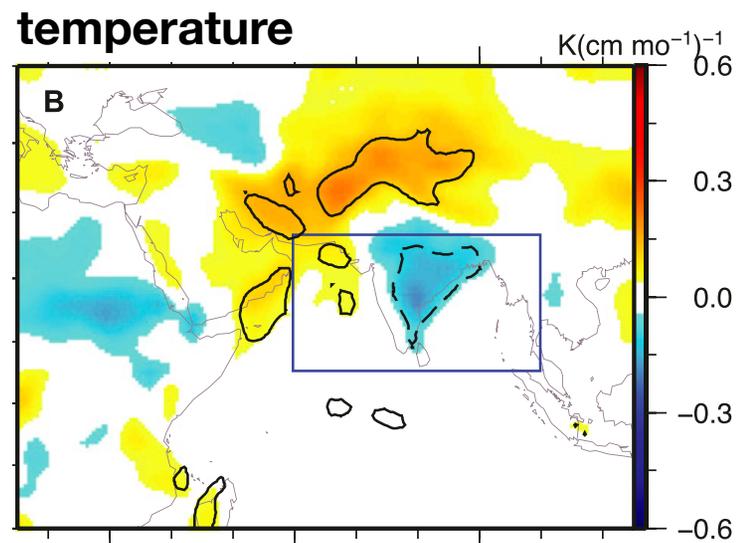
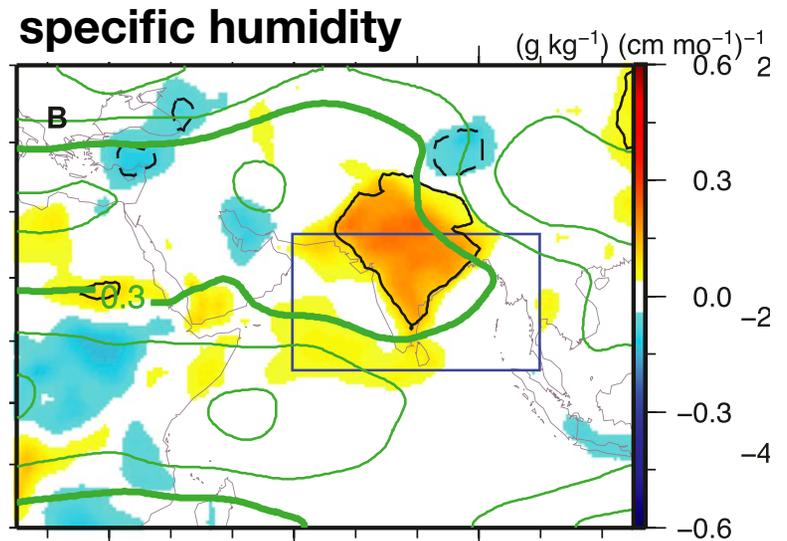
Convective quasi-equilibrium framework is also consistent with South Asian interannual variability

regressions on South Asian monsoon rainfall of near-surface ...



Data: GHCN rainfall & ERA-Interim
T, q, moist static energy

Hurley & Boos (2013)



How to think about monsoon circulations?

i.e. given a forcing, how will the precipitating circulation respond?

Dominant categories of theory:

- Surface temperature gradients drive flow **wrong**

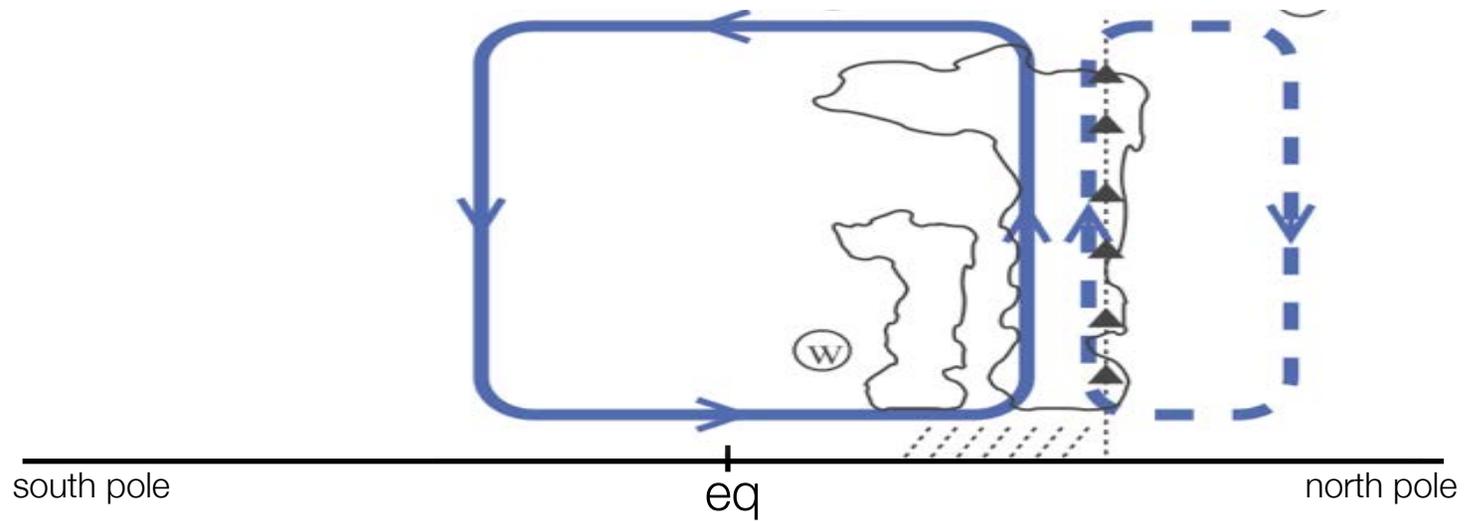
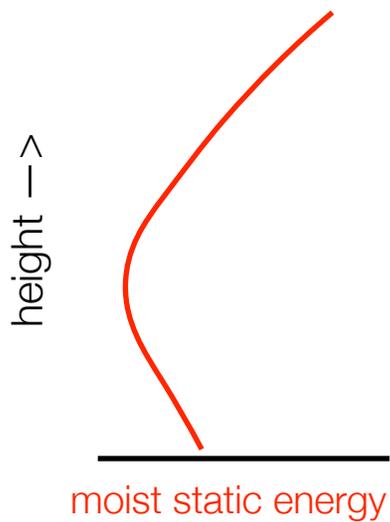
- Convective quasi-equilibrium **mechanistic but diagnostic**

- Vertically integrated atmospheric energy budget **prognostic but not mechanistic**

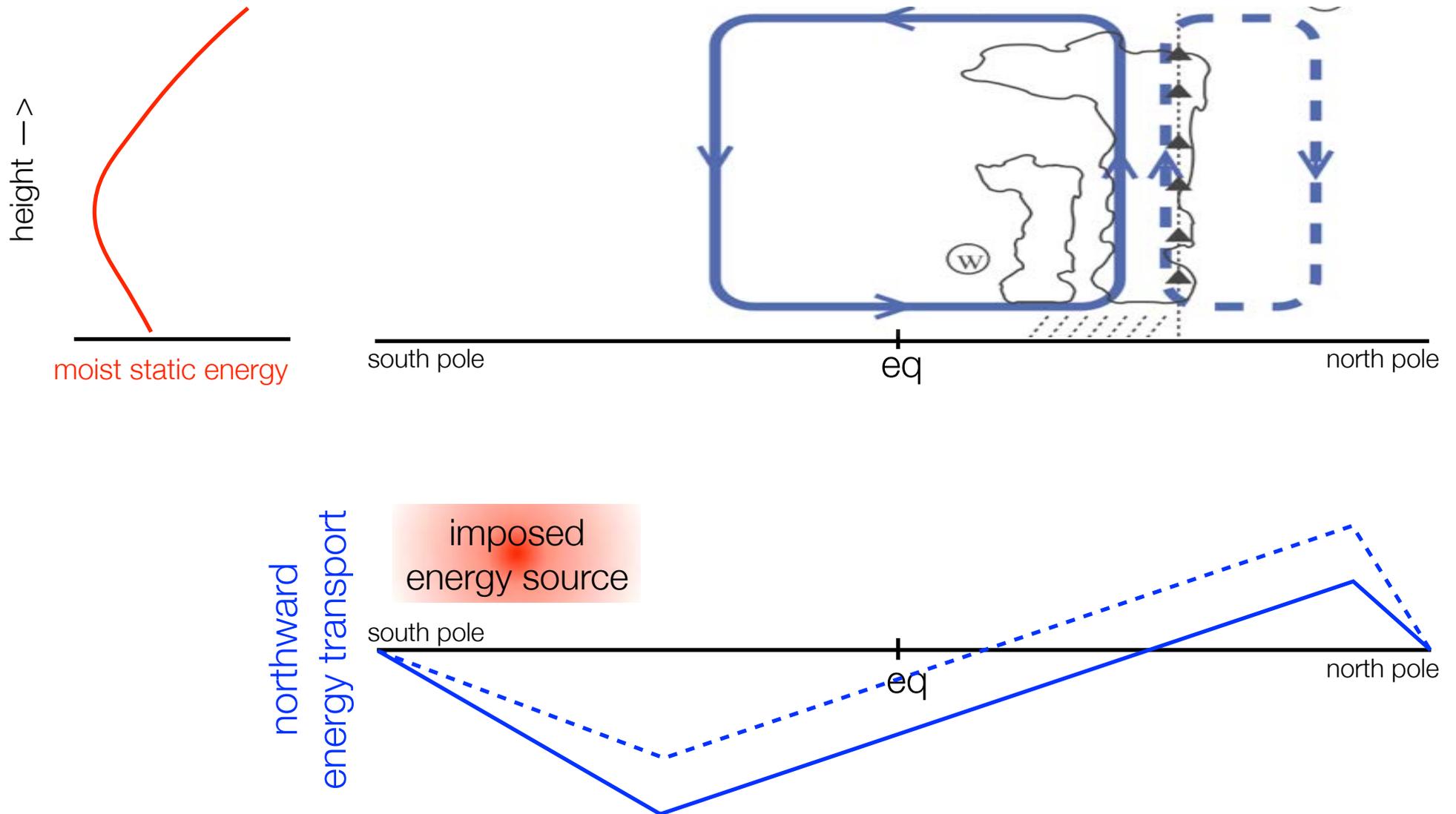
Atmospheric meridional overturning circulations transport energy

$$\nabla \cdot \int_{p_s}^0 v h \frac{dp}{g} = Q$$

energy input to atmospheric column

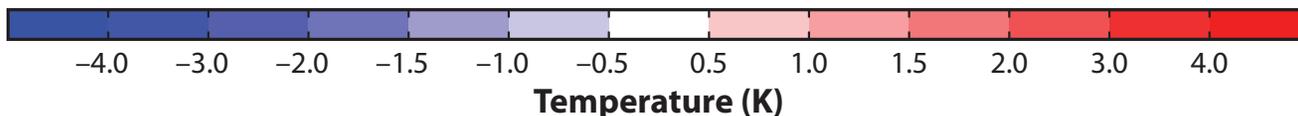
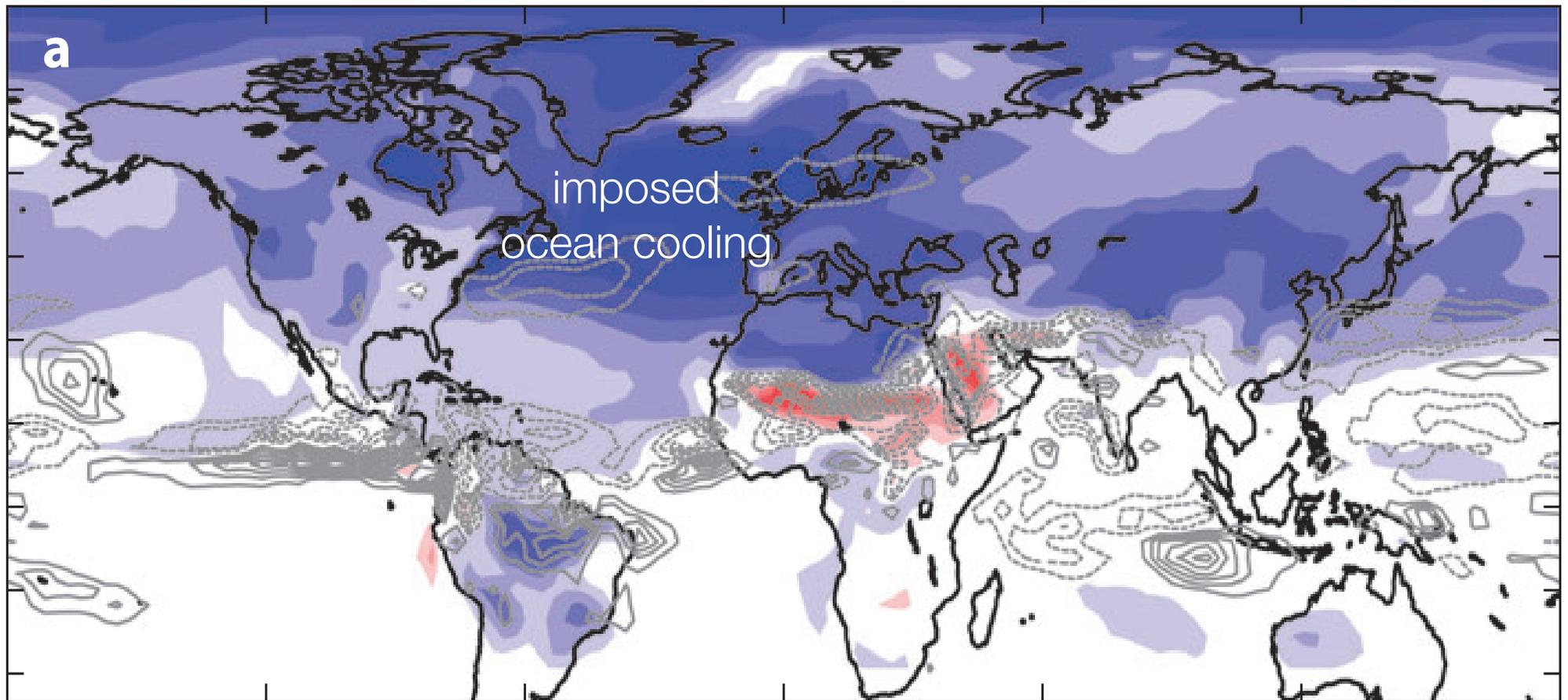


Precipitation peak shifts in response to an imposed high latitude energy source



GCMs & paleo proxies show that high-latitude thermal forcings can produce ITCZ shifts

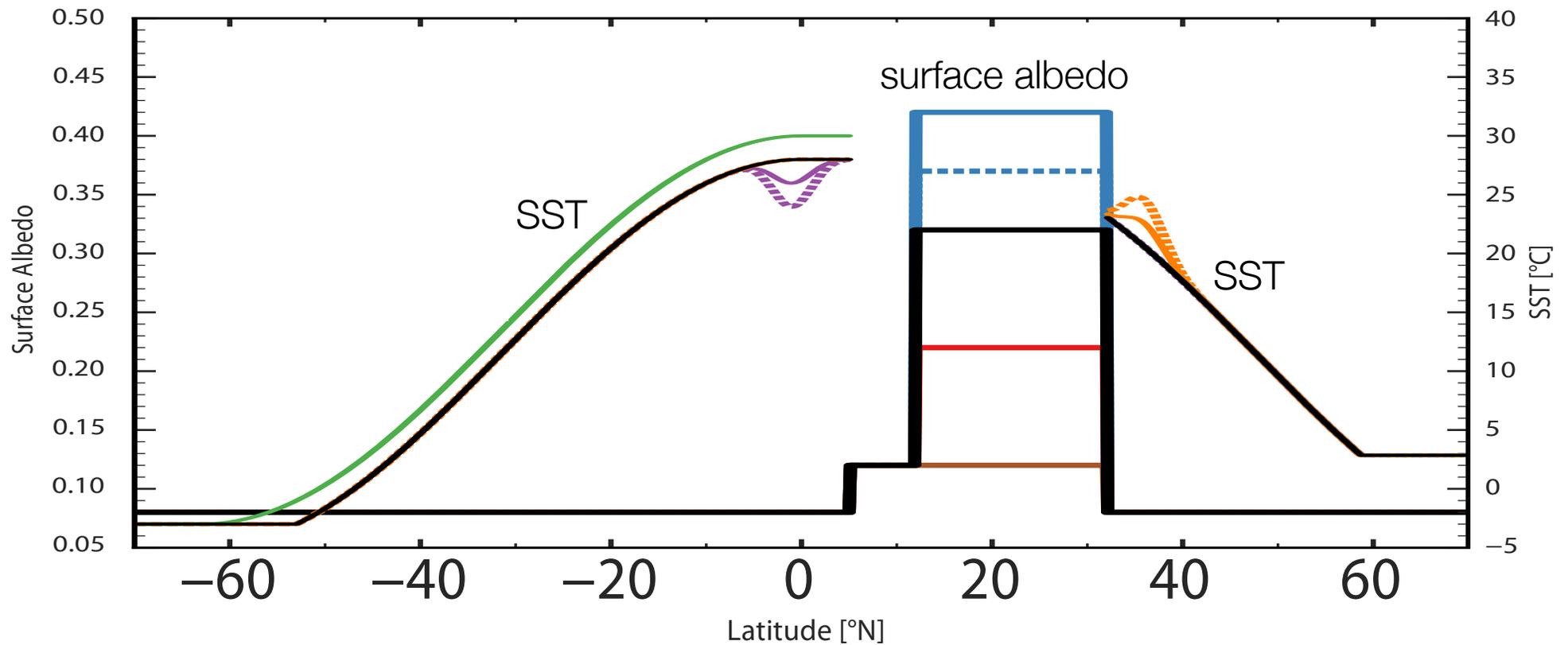
colors = surface temperature anomaly, contours = precipitation anomaly



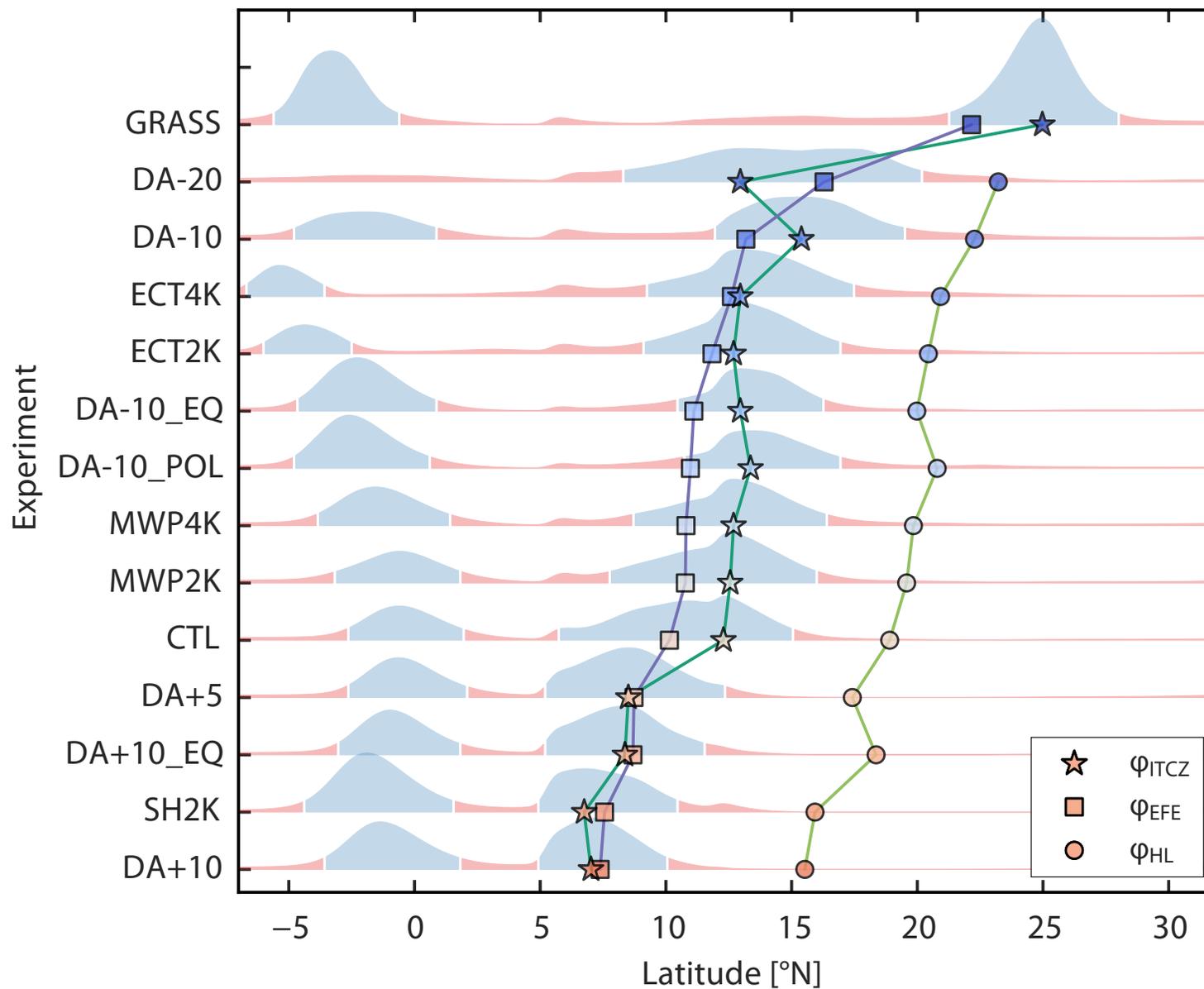
Chiang & Friedman (2012)

This works quantitatively for a variety of forcings in an idealized beta-plane model

imposed land surface albedo and SST boundary conditions in WRF model

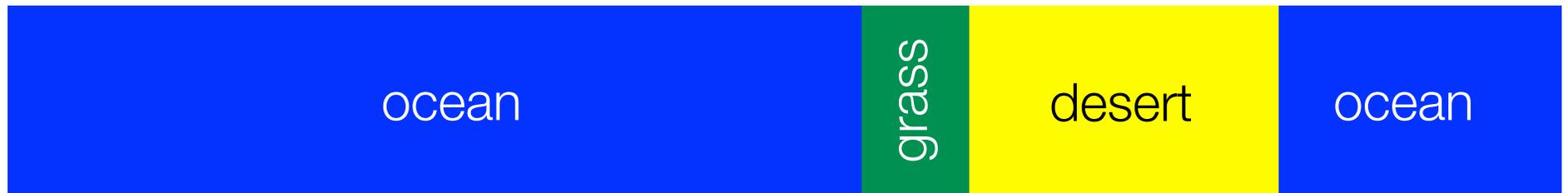
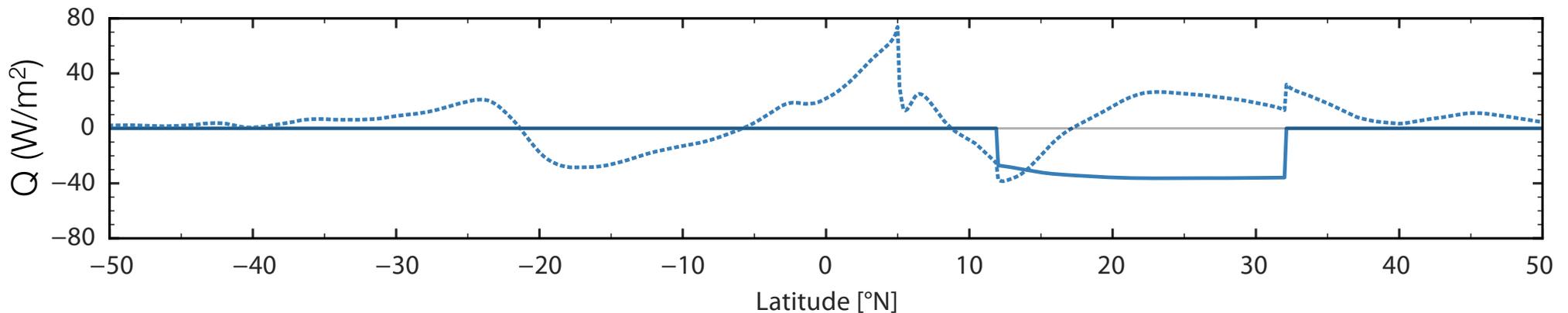


Precipitation maximum and energy flux equator generally shift together



But energetic feedbacks are just as large as the imposed forcing

forcing (solid) and feedback (dashed) for increase in surface albedo of desert



Summary

- Convective quasi-equilibrium and the vertically integrated atmospheric energy budget can both be used to understand the response of monsoons to forcings:
 - interannual variability (enhanced land-ocean h contrast = more rain)
 - ITCZ shifts toward an applied energy source, but there are large energetic feedbacks
- An energetically closed simple model & a GCM show that monsoon strength exhibits no threshold response to a broad range of forcings